

冀北承德甲山正长岩——燕山陆内造山带 岩石圈减薄的早期记录

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内容提要 早白垩世承德甲山正长岩体位于燕山陆内造山带东段,承德逆掩片东侧,可划分为三个单元,从早到晚,依次为姜家湾单元、龙潭沟单元和龙潭南沟单元。化学成分上,该岩体Si、全碱、TFe、REE、Th、Ga、Nb、Zr、Hf含量较高,Mg、Ba、Sr、Ti、Cr、Co、Ni、V等过渡元素亏损,Ga/Al值大,Ce/Pb值在6.12~13.41,富轻稀土,中等铕负异常,具有A型花岗岩的特点。岩体形成于造山晚期岩石圈伸展的环境下,与地幔上涌有关,说明区域晚侏罗世的挤压缩短和地壳增厚转变为早白垩世的伸展和岩石圈减薄,暗示中国东部的岩石圈减薄应从早白垩世开始。

关键词 正长岩 地球化学特征 构造意义 岩石圈减薄 承德 河北

燕山造山带位于华北板块北缘,中生代陆内造山强烈(赵越,1990; Chen, 1998; 郑亚东等,2000; Davis et al., 2001),岩浆活动异常发育,近年来受到地质界的关注(赵越,1990; 白志民等,1991,1999; 牟保磊等,1992; 刘凤山等,1998; 阎国翰等,1998; 朱大岗等,1999; 许保良等,1999a, 1999b; 邓晋福等,2000; 张旗等,2001a, 2001b; 邵济安等,1999, 2001; 李伍平等,2001a, 2001b; 钱青等,2002)。对于岩浆的成因学者间意见并不一致,有的研究者认为是太平洋板块与亚洲大陆相互作用形成的太平洋型活动陆缘(任纪舜等,1980),岩浆活动产于挤压构造环境下(程裕淇,1994);有人提出是太平洋板块低角度俯冲的结果(Yoshio et al., 1994; 赵海玲等,1998);邓晋福等(1996)提出华北克拉通岩石圈与伊泽

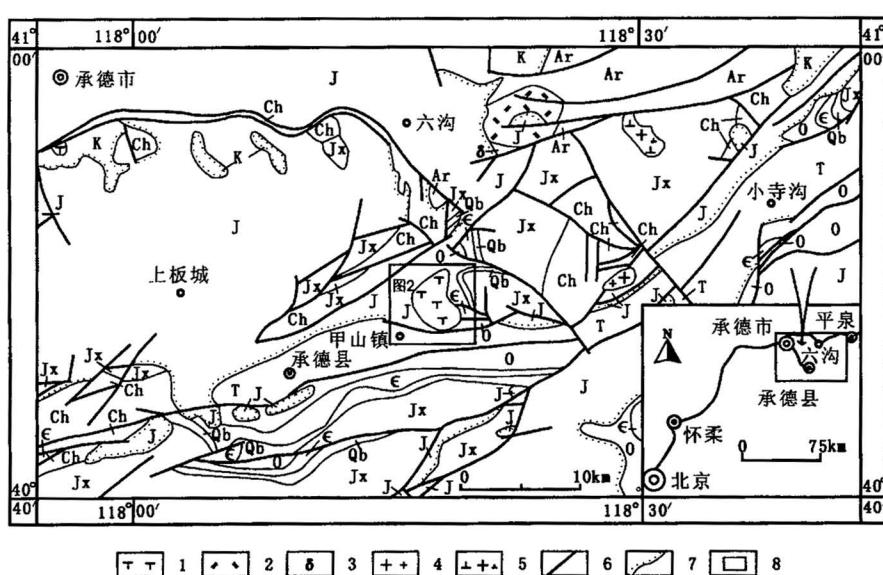


图1 承德甲山区域地质图(据河北省北京市天津市区域地质志,1989,简编)
Fig. 1 Regional geological map of Jiashan area, Chengde (after Regional Geology of Hebei Provinces, Beijing and Tianjin Municipality, 1989)

K—白垩系; J—侏罗系; T—三叠系; O—奥陶系; ←—寒武系; Qb—青白口系; Jx—蓟县系; Ch—长城系; Ar—太古界; 1—正长岩; 2—二长岩; 3—闪长岩; 4—花岗斑岩; 5—花岗闪长斑岩; 6—断层; 7—不整合界线; 8—研究区
K—Cretaceous; J—Jurassic; T—Triassic; O—Ordovician; ←—Cambrian; Qb—Qingbaikou System; Jx—Jixian System; Ch—Changcheng System; Ar—Archaeon; 1—syenite; 2—monzonite; 3—diorite; 4—granite—porphyry; 5—grano—diorite—porphyry; 6—fault; 7—unconformity boundary; 8—study area

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奈崎大洋岩石圈的汇聚形成的弧火山岩;多数地质学家(Xu, 1990; 赵越等, 1994; 朱大岗等, 1999; 郑亚东等, 2000)将燕山构造—岩浆带形成与太平洋—欧亚板块的相互作用相联系;邵济安等(2001)、李伍平等(2001a, 2001b)和张旗等(2001b)则认为燕山地区中生代岩浆的形成和演化与板内的底侵作用有关,而与太平洋板块俯冲无直接的成因联系。

该区是燕山陆内造山带代表性的晚侏罗世地壳挤压增厚的地区,在侏罗纪燕山地区以挤压缩短和地壳增厚为主,白垩纪转变为区域伸展和岩石圈减薄,什么时间由挤压转变为伸展?其证据是什么?是地质界广泛关注的问题,甲山岩体的研究将对区域地壳和岩石圈减薄过程提供重要的资料。笔者曾在1997年论述过甲山岩体各单元划分和地球化学特征,本文则以甲山岩体的详细调查和系统采样为基础,通过岩体的岩相学、矿物学和地球化学特征,阐明区域早白垩世的岩石圈减薄。

1 地质背景

承德甲山岩体位于燕山陆内造山带的东段。区域上侏罗系较发育(图1),为一套火山熔岩、火山碎屑岩和沉积岩,花岗岩较发育。区域上白垩纪早期发育一套安山岩—粗安岩—粗面岩组合的钙碱—碱性系列的中性熔岩(朱大岗等, 1999),晚期侵入岩较发育,以酸性岩(如石洞子沟花岗斑岩)和碱性岩为主,岩石组合为闪长玢岩、花岗斑岩和正长岩,成因类型以I型和A型为主,燕辽地区晚侏罗世—早白垩世的中酸性火山岩和侵入岩具有埃达克岩的特征(张旗等, 2001a, 2001b, 2001c)。

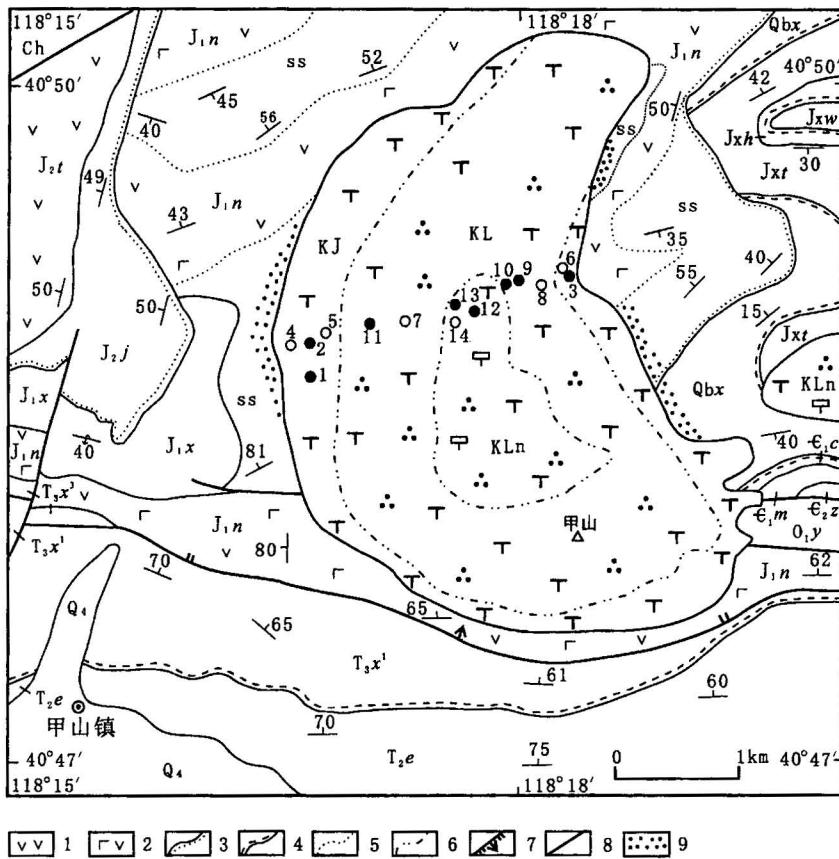


图2 承德甲山正长岩地质图(根据1:5万上谷幅区域地质图改编)

Fig. 2 Geological map of Jiashan syenite, Chengde

(after the 1:50000 Shanggu station regional geological map)

Q_4 —第四系全新统; J_{2t} —中侏罗统髻髻山组; J_{2j} —中侏罗统九龙山组; J_{1x} —下侏罗统下花园组; J_{1n} —下侏罗统南大岭组; ss —下侏罗统南大岭组中的砂岩、粉砂岩、页岩;下三叠统杏石口组; T_{3x^3} —三段, T_{3x^1} —一段; T_{2e} —中三叠统二马营组; O_{1y} —下奥陶统冶里组; \leftarrow_{2z} —中寒武统张夏组; \leftarrow_{1m} —下寒武统馒头组; \leftarrow_{1c} —下寒武统昌平组; Qbx —青白口系下马岭组;

Jxt —蓟县系铁岭组; Jxh —蓟县系洪水庄组; Jxw —蓟县系雾迷山组; Ch —长城系; KLn —龙潭南沟单元; KL —龙潭沟单元; KJ —姜家湾单元; 1—安山岩; 2—玄武安山岩; 3—角度不整合界线; 4—平行不整合界线; 5—岩性界线; 6—涌动型侵入接触界线; 7—正断层; 8—性质不明断层; 9—角岩化; ○—1997年采样点, 4代表表1样品序号; ●—本次采样点, 1代表表1样品序号; △—山峰

Q_4 —Quaternary; J_{2t} —Middle Jurassic Tiaojishan Formation; J_{2j} —Middle Jurassic Jiulongshan Formation; J_{1x} —Lower Jurassic Xiahuyuan Formation; J_{1n} —Lower Jurassic Nandaling Formation; ss —sandstone, siltstone, shale of Lower Jurassic Nandaling Formation; T_{3x^3} —Third Member of Lower Triassic Xingshikou Formation; T_{3x^1} —First Member of Lower Triassic Xingshikou Formation; T_{2e} —Middle Triassic Ermaying Formation; O_{1y} —Lower Ordovician Yeli Formation; \leftarrow_{2z} —Middle Cambrian Zhangxia Formation; \leftarrow_{1m} —Lower Cambrian Mantou Formation; \leftarrow_{1c} —Lower Cambrian Changping Formation; Qbx —Xiamaling Formation of Qingbaikou System; Jxt —Tieling Formation of Jixian System; Jxh —Hongshuizhuzhuang Formation of Jixian System; Jxw —Wumishan Formation of Jixian System; Ch —Changcheng System; KLn —Longtannangou unit; KL —Longtangou unit; KJ —Jiangjiawan unit; 1—andesite; 2—basaltic andesite; 3—angular unconformity boundary; 4—parallel unconformity boundary; 5—lithologic boundary; 6—salient intrusive contact boundary; 7—normal fault; 8—fault of uncertain nature; 9—hornfels; ○—location of the samples taken in 1997, 4 represents sample number in table 1, ●—location of the samples for this paper, 1 represents sample number in table 1; △—peak

岩体出露于河北省承德县甲山镇附近,构造上位于平泉—古北口断裂(赵越,1990)南侧,但紧邻 Davis 等(1998)“未名逆冲断层”。岩体南侧为下板城盆地,发育三叠系到上侏罗统,在晚侏罗世强烈褶皱变形。西侧紧邻 Davis 等(1998)所确定的逆冲岩片,其逆冲位移至少 40~45km(郑亚东等,2000),逆冲断层的时代为晚侏罗世(161~132Ma),虽然对该逆冲岩片还有不同认识,但晚侏罗世燕山地区地壳挤压增厚作用是普遍存在的(和政军等,1998;杨庚等,2001;张长厚等,2002)。岩体南北向长 5 km,最宽东西向 3.8 km,平面上为不规则椭圆形,出露面积约 14 km²(图 2),岩体侵入最新地层为下侏罗统南大岭组(孙跃武等,1996;杨富全等,1999),并侵入承德县逆冲断层中。岩体 K-Ar 全岩年龄 94.94 ± 2.25 Ma(杨富全等,1997);锆石 U-Pb 年龄 113 ± 2 Ma(郑亚东等,2000;Davis et al., 2001),时代为早白垩世晚期。围岩热接触变质明显,尤其是岩体西边与南大岭组地层接触带上,外接触带从西向东发育角岩化→含红柱石长英质角岩→红柱石长英质角岩,围岩中硅化也很发育。岩体内接触带发育冷凝带。岩体与围岩有规则而清楚的接触界线,接触面平直外倾。按岩石谱系单位划分原则将甲山岩体进行解体,圈定 4 个侵入体,划分了 3 个单元。除一个侵入体分布在主岩体东部围岩中,其他 3 个侵入体,在时间上密切相关,在空间上为套叠式同心环带状分布,共同组成一个同心分布的同源岩浆侵入序列。

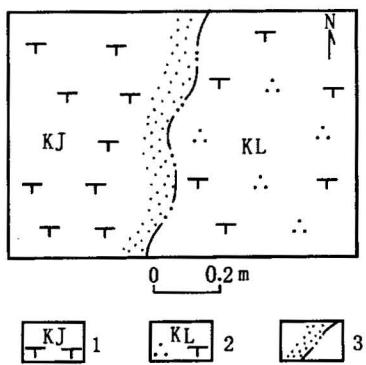


图 3 承德甲山岩体龙潭沟单元涌动侵入姜家湾单元接触关系素描图

Fig. 3 Sketch showing the contact relation that Longtangou unit salient intrude Jiangjiawan unit of the Jiashan intrusive rocks, Chengde
1—姜家湾单元; 2—龙潭沟单元;
3—涌动型侵入接触界线及过渡带
1—Jiangjiawan unit; 2—Longtangou unit;
3—salient intrude contact boundary and intermediate belt

3 个单元从早到晚依次为姜家湾单元,分布于岩体最外侧;龙潭沟单元分布在岩体内侧,出露面积最大;龙潭南沟单元出露于岩体中心。龙潭沟单元与姜家湾单元之间没有截然的侵入界线,但在十几厘米范围内岩石矿物成分及结构快速变化,为涌动型侵入接触关系(图 3),龙潭南沟单元与龙潭沟单元呈涌动型侵入接触关系,二者之间有清晰的岩性界面,成分上发生突变,由不含包体到含大量包体,结构上由粗粒结构快速变为似斑状结构,暗色矿物条带、长石斑晶平行接触面分布,指示龙潭南沟单元侵入时间稍晚(图 4)。

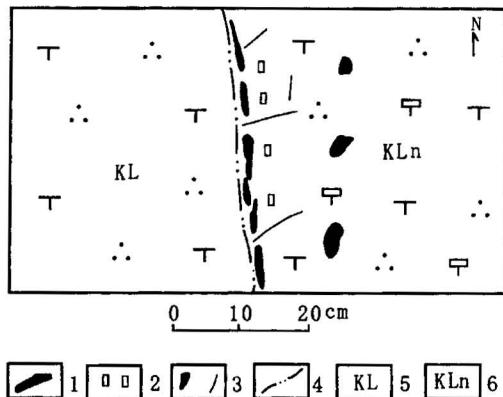


图 4 承德甲山龙潭南沟单元涌动侵入龙潭沟单元接触关系素描图

Fig. 4 Sketch showing the contact relation that Longtannangou unit salient intrude Longtangou unit of the Jiashan intrusive rocks, Chengde
1—暗色矿物带; 2—长石斑晶; 3—包体和节理; 4—涌动型侵入接触界线; 5—龙潭沟单元; 6—龙潭南沟单元
1—Dark mineral belt; 2—feldspar porphyritic crystal;
3—enclosure, joint; 4—alient intrude contact boundary;
5—Longtangou unit; 6—Longtannangou unit

2 岩石学特征

姜家湾单元: 岩性为粗中粒含角闪辉石正长岩,粗中粒半自形粒状结构,局部交代残留结构。主要矿物成分: 钾长石(条纹长石和正长石)约占 70%~80%, 半自形—它形粒状, 粒径一般 2~5mm, 部分 5~7mm; 普通角闪石(5%~10%); 单斜辉石(5%~10%), 半自形柱粒状, 粒径 0.1~2.7mm; 斜长石(5%)、石英(<5%)。副矿物组合: 磁铁矿、锆石、磷灰石、钛铁矿、黄铁矿、辉钼矿, 见方铅矿、闪锌矿。

龙潭沟单元: 岩性为粗粒辉石角闪正长岩,粗粒半自形粒状结构,块状构造。矿物成分: 条纹长

石(62%~85%),半自形—它形粒状,粒径一般5~8mm,部分2~5mm;普通角闪石(5%~10%),柱状,粒度0.5~2mm;单斜辉石(4%~7%),粒径0.1~1.5mm;斜长石(2%~5%),石英(5%~8%)。副矿物主要为磁铁矿、褐铁矿、锆石、磷灰石、黄铁矿、黄铜矿,见辉钼矿、方铅矿、钛铁矿、磷钇矿和孔雀石。

龙潭南沟单元:岩性为中细粒似斑状含辉石角闪石英正长岩,块状构造,似斑状结构,斑晶(30%)以正长石为主,粒径为2.5~15mm,少量斜长石和单斜辉石。基质具中细粒半自形粒状结构,主要矿物成分:正长石(45%~50%)、角闪石(5%~10%)、石英(5%~7%)、单斜辉石(3%~5%)、斜长石(5%)和少量黑云母。副矿物主要为磁铁矿、磷灰石、锆石、黄铁矿、钛铁矿,少量榍石、磷钇矿、辉钼矿、方铅矿、闪锌矿、褐铁矿。该单元中含同源包体,其形态以圆状、椭圆状居多,另有长方形和不规则状,大小一般在1~3cm,较大者7~20cm,最大者达30cm,较大包体中含有长石斑晶。包体岩性为角闪辉石石英二长闪长岩,似斑状结构,斑晶为正长石(5%~10%)和少量辉石。基质为细粒结构,由钾长石(10%~15%)、斜长石(45%~50%)、石英(10%~15%)、辉石(10%~15%)和角闪石(5%~10%)组成。副矿物为磁铁矿、磷灰石和锆石。

3 地球化学特征

甲山岩体虽然可划分出三个单元,但各单元之间成分差异不是很大,属基本均匀的岩体,因此,选择穿过岩体中间的剖面采样。样品新鲜,由中国地质科学院地球物理地球化学勘查研究所用无污染法破碎、磨碎(>200目)制成分析样品。主量元素FeO采用容量法,CO₂、Corg采用电导法,H₂O⁺采用重量法,其他用熔片法X-射线荧光光谱法(XRF)分析;稀土元素采用等离子体质谱法(ICP-MS)分析,检出限为0.1×10⁻⁶(Ce为5.0×10⁻⁶);微量元素采用等离子体质谱法和X荧光光谱法分析。分析过程中用GSR1、GRS2、GRS5监控,并对部分样品进行重复测定,一级标准物质合格率100%,重复样

合格率100%。

3.1 岩石化学特征

表1列出岩体三个单元代表性样品的主要元素分析结果。各单元岩石成分具有一定的差异,姜家湾单元SiO₂含量变化于61.34%~63.15%,平均为62.47%;K₂O+Na₂O=10.44%~11.56%,平均为11.03%;里特曼指数*n*变化于5.70~6.2。龙潭沟单元SiO₂含量介于63.62%~66.46%,平均为65.27%,K₂O+Na₂O=10.26%~11.15%,平均10.60%,里特曼指数*n*介于4.53~5.60。龙潭南沟单元SiO₂含量变化于64.77%~65.85%,平均

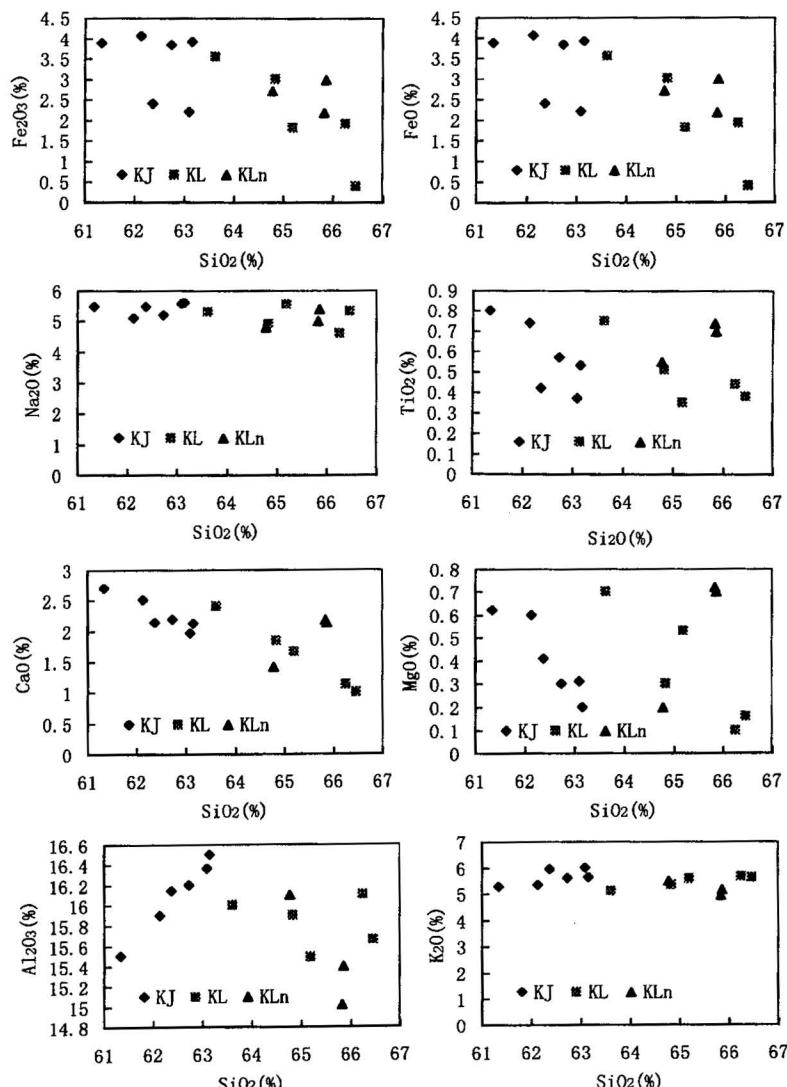


图5 承德甲山正长岩主量元素的Harker图解

Fig. 5 Harker diagrams of major elements for the

Jiashan syenite, Chengde

◆KJ—姜家湾单元;■KL—龙潭沟单元;▲KLn—龙潭南沟单元
◆KJ—Jiangjiawan unit; ■KL—Longtangou unit; ▲KLn—Longtannangou unit

表 1 承德甲山正长岩各单元主要元素(%)分析
Table 1 Major elements (%) of each unit from Jiashan syenite, Chengde

序号	单元号	样号	SiO ₂	TiO ₂	Al ₂ O ₃	FeO	Fe ₂ O ₃	MnO	CaO	MgO	K ₂ O	Na ₂ O	P ₂ O ₅	H ₂ O ⁺	H ₂ O ⁻	Los	CO ₂	Corg	总量	A/CNK	6
1	KJ	D260	62.73	0.57	16.2	3.84	2.13	0.17	2.19	0.3	5.60	5.2	0.14	0.48		0.18	0.09	99.82	0.87	5.91	
2	KJ	D272	63.15	0.53	16.5	3.92	2.00	0.17	2.12	0.2	5.64	5.6	0.13	0.55		0.11	0.07	100.69	0.87	6.27	
3	KJ	D277	62.13	0.74	15.9	4.06	2.54	0.18	2.51	0.6	5.34	5.1	0.2	0.66		0.66	0.12	100.74	0.85	5.70	
4	KJ	YQ060	63.09	0.37	16.36	2.21	3.48	0.14	1.96	0.31	5.99	5.57	0.12		0.16	0.71		100.50	0.85	6.65	
5	KJ	YQ061	62.37	0.42	16.14	2.40	3.64	0.18	2.14	0.41	5.94	5.47	0.12		0.09	0.64		99.96	0.83	6.72	
6	KJ	YQ067	61.34	0.8	15.5	3.88	3.16	0.20	2.7	0.62	5.27	5.47	0.10		0.05	1.02		100.11	0.79	6.29	
7	KL	YQ063	66.46	0.38	15.66	0.40	3.98	0.12	1.00	0.16	5.64	5.32	0.12		0.06	1.02		100.32	0.94	5.12	
8	KL	YQ066	65.19	0.35	15.49	1.82	2.7	0.12	1.66	0.53	5.59	5.56	0.09		0.11	0.96		100.88	0.85	5.60	
9	KL	D262	64.83	0.51	15.9	3.01	1.64	0.12	1.84	0.3	5.35	4.9	0.12	0.30		0.73	0.07	99.62	0.93	4.81	
10	KL	D276	63.62	0.75	16.0	3.56	1.52	0.13	2.41	0.7	5.10	5.3	0.21	0.68		0.73	0.05	100.70	0.86	5.25	
11	KL	D273	66.25	0.44	16.1	1.92	2.42	0.11	1.13	0.1	5.66	4.6	0.09	0.59		0.33	0.03	99.77	0.98	4.53	
12	KLn	D261	65.85	0.70	15.4	3.00	1.45	0.11	2.14	0.7	5.18	5.4	0.19	0.44		0.07	0.03	100.70	0.84	4.90	
13	KLn	D275	64.77	0.55	16.1	2.72	2.37	0.12	1.42	0.2	5.52	4.8	0.14	0.36		0.33	0.14	99.54	0.99	4.89	
14	KLn	YQ065	65.82	0.74	15.02	2.19	2.98	0.12	2.18	0.72	4.95	5.00	0.11		0.07	0.41		100.31	0.85	4.34	

注:序号 4~8 和 14 据杨富全等,1997;其余由中国地质科学院地球物理地球化学勘查研究所用熔片法 X-射线荧光光谱法测试。

65.48%, K₂O + Na₂O = 9.95%~10.5%, 平均 10.28%, 里特曼指数介于 4.34~4.90。碱性指数(AKI)值变化于 0.85~0.98;多数样品 K₂O > Na₂O, Na₂O/K₂O = 0.81~1.04, 多在 0.92~0.99

之间。在哈克图解中(图 5), SiO₂ 与 MgO、TiO₂、CaO、Na₂O、FeO、Fe₂O₃ 呈负相关, 而与 Al₂O₃ 则出现先增后减的趋势, 表明岩浆经历了结晶分异。在 K₂O—Na₂O 图解(Collins et al., 1982; 图略)上三个

表 2 承德甲山正长岩各单元稀土(μg/g)分析结果

Table 2 Rare earth elements compositions (μg/g) of each unit from Jiashan syenite, Chengde

序号	单元	样号	La	Ce	Pr	Nd	Sm	Eu	Gd	Tb	Dy	Ho	Er	Tm	Yb	Lu
1	KJ	D260	53	92	12.3	45.8	8.6	1.85	7.7	1.21	6.3	1.29	3.2	0.60	3.4	0.61
2	KJ	D272	68	120	15.5	57.6	10.6	2.18	9.2	1.48	7.5	1.51	3.9	0.70	4.0	0.68
3	KJ	D277	72	140	17.3	65.3	12.2	2.47	10.8	1.74	8.8	1.88	4.7	0.81	4.7	0.75
4	KJ	P50XT27	67.9	114.7	11.95	50.98	8.57	1.97	8.55	1.31	6.88	1.45	4.07	0.57	3.7	0.49
5	KL	D262	108	185	21.4	73.1	12.5	1.72	11.1	1.77	9.5	1.97	5.5	1.00	5.7	0.95
6	KL	D276	73	134	15.4	55.8	9.8	2.03	8.8	1.41	7.3	1.53	4.1	0.71	4.1	0.67
7	KL	D273	115	175	23.8	83.4	14.0	1.55	12	1.88	9.6	1.96	5.2	0.91	5.2	0.88
8	KL	P50XT11	75.19	138	15.45	67.02	12.2	2.18	11.2	2.0	10.12	2.11	5.78	0.76	4.85	0.81
9	KLn	D261	96	148	19.5	67	11.3	1.68	10.0	1.63	8.5	1.79	4.9	0.88	4.9	0.80
10	KLn	D275	78	142	16.7	59.4	10.8	1.87	9.3	1.49	8.0	1.71	4.4	0.81	4.5	0.75
11	KLn	P50XT21	94.43	158.4	15.61	62.15	10.01	1.51	9.52	1.73	8.16	1.77	5.06	0.70	4.51	0.63

序号	单元	样号	LREE	HREE	LREE/HREE	ΣREE	(La/Sm) _N	(La/Yb) _N	δEu
1	KJ	D260	213.6	24.31	8.78	237.9	3.88	10.51	0.68
2	KJ	D272	273.88	28.97	9.45	302.85	4.04	11.46	0.66
3	KJ	D277	309.27	34.18	9.05	343.45	3.71	10.33	0.64
4	KJ	P50XT27	256.07	27.02	9.48	283.09	4.99	12.37	0.70
5	KL	D262	401.7	37.49	10.72	439.2	5.44	12.78	0.44
6	KL	D276	290.03	28.62	10.13	318.65	4.69	12	0.66
7	KL	D273	412.8	37.63	10.97	450.4	5.17	14.91	0.36
8	KL	P50XT11	310.04	37.63	8.24	347.67	3.88	10.54	0.56
9	KLn	D261	343.48	33.4	10.28	376.88	5.34	13.21	0.47
10	KLn	D275	308.77	30.96	9.97	339.73	4.54	11.69	0.56
11	KLn	P50XT21	342.11	32.08	10.66	374.19	5.94	14.10	0.47

注: 样品由中国地质科学院地球物理地球化学勘查研究所采用等离子体质谱法(ICP-MS)分析。

单元均位于A型花岗岩区;铝的含量高, $\text{Al}_2\text{O}_3 = 15.02\% \sim 16.5\%$, 岩石的铝饱和指数 A/CNK 值介于 $0.79 \sim 0.99$, 为准铝质;碱度率(AR)较高, 变化于 $3.60 \sim 4.85$, 在 AR— SiO_2 图解(Wright, 1969; 图略)中全部位于碱性岩区, 表现岩石属碱性系列范围。

综上所述, 甲山岩体在化学成分上具富硅、富碱、富铁、富铝、贫钛、贫镁的特点, 从早单元到晚单元, SiO_2 含量增加, $\text{K}_2\text{O} + \text{Na}_2\text{O}$ 则逐渐减少, 里特曼指数 δ 明显减少, 表明三个单元在成分上具亲缘关系的演化序列, 从碱性向酸碱性演化。

3.2 稀土元素特征

甲山岩体各单元稀土元素数据列于表2。由表2知:甲山岩体最显著的特征是稀土元素总量较高, ΣREE 为 $237.86 \sim 450.38\mu\text{g/g}$; LREE/HREE 为 $6.05 \sim 10.97$, $n_{\text{La}}/n_{\text{Sm}} = 3.71 \sim 5.94$, $n_{\text{La}}/n_{\text{Yb}} = 10.33 \sim 14.91$, 轻重稀土分馏强烈, 轻稀土明显富集, 具有中等明显的负铕异常($\delta\text{Eu} = 0.36 \sim 0.70$), 显示岩浆曾发生过斜长石结晶分异(邱检生等, 2000; 李伍平等, 2001c)。在稀土配分图上显示出右倾, 且 LREE 富集和 Eu 负异常的特征(图6)。各单元稀土特征呈规律性变化, 从早单元到晚单元稀土总量增高再降低(平均值 $291.81\mu\text{g/g} \rightarrow 388.99\mu\text{g/g} \rightarrow 363.61\mu\text{g/g}$); δEu 值从早单元 → 晚单元降低(δEu 平均值 $0.67 \rightarrow 0.51 \rightarrow 0.5$), 镐负异常愈来愈明显。三个单元稀土元素配分曲线相似, 稀土元素之间的分异接近, 且集中在一个很狭小的区域内, 反映它们具有相同的源区性质(匡少平等, 2001)。

3.3 微量元素特征

微量元素(表3)显示, 甲山岩体的 Ba、Sr 和 Cr、Co、Ni、V 等过渡元素亏损, 而高场强元素(HFSE)

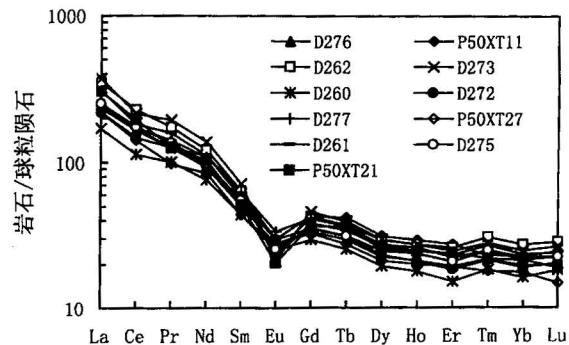


图6 承德甲山岩体稀土元素配分模式
Fig. 6 Chondrite-normalized REE patterns of the Jiashan syenite, Chengde

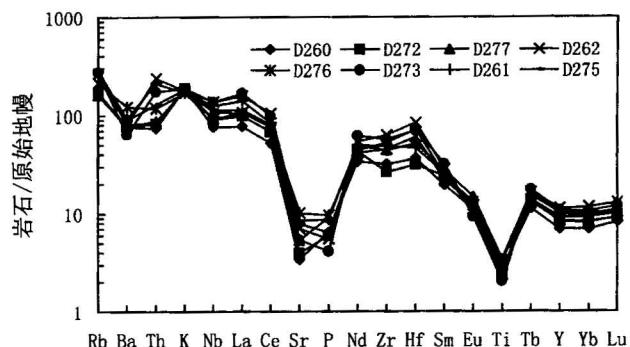


图7 承德甲山正长岩微量元素蛛网图
(据 Sun 等, 1989)

Fig. 7 Primitive mantle-normalized trace element spidergrams for Jiashan syenite, Chengde
(after Sun et al., 1989)

如 Nb、Zr、Hf 含量较高, 具有 A型花岗岩的特征(Whalen et al., 1996)。 $\text{Ce}/\text{Pb} = 6.12 \sim 13.41$, 具有裂谷特征。在原始地幔标准化微量元素蛛网图上(图7)显示三个单元曲线形态相同, 并且 Ba、Sr、P、Ti 呈明显的“V”型谷。岩体的大离子亲石元素 Rb、Th 富

表3 承德甲山正长岩微量元素($\mu\text{g/g}$)

Table 3 Trace elements($\mu\text{g/g}$) of the Jiashan syenite, Chengde

序号	单元	样号	Sr	Rb	Ba	Th	U	Cr	Ta	Nb	Zr	Hf	Cu	Pb	Zn	Mo	V	Ni	Y	Co	Ga	Sc
1	KJ	D260	72	102	533	6.3	1.22	5.0	3.63	54.2	353	11.2	6.5	8.5	120	2.71	19.9	5.6	31.5	1.6	30.2	6.4
2	KJ	D272	83	108	537	7.0	1.48	5.3	4.24	64.5	295	9.7	6.7	10.7	117	2.95	16.3	7.9	37.5	1.8	30.7	6.9
3	KJ	D277	112	102	555	7.4	1.52	5.0	4.93	77.5	493	15.8	9.5	12.8	127	3.22	25.3	6.6	45.6	3.9	30	8.7
4	KL	D262	148	170	563	20.3	4.88	5.0	7.52	97.9	701	25.8	5.3	13.8	107	4.17	24.0	5.9	51.1	2.1	29.8	4.6
5	KL	D276	214	136	855	10.0	2.59	6.2	4.58	65.9	502	18.4	8.1	15.9	118	3.76	26.5	5.8	38.5	3.9	27.8	5.6
6	KL	D273	114	174	447	14.8	2.11	5.0	6.65	97.4	633	21.7	7.6	17.8	113	2.93	18.9	6.1	48.7	1.7	29.9	4.1
7	KLn	D261	180	178	647	18.0	3.53	5.0	6.60	89.5	537	14.3	9.5	24.2	100	4.01	29.3	5.7	46.4	4.8	25.7	6.0
8	KLn	D275	167	149	631	10.9	1.70	5.0	5.31	80.2	585	21.9	5.7	13.3	98	1.93	8.9	2.9	41.5	2.6	28.6	4.6

注:样品由中国地质科学院地球物理地球化学勘查研究所测试;Sr、Rb、Ba、Cr、Nb、Zn、V、Ni、Ga 元素采用 XRF 分析, 其他元素采用 ICP-MS 分析。

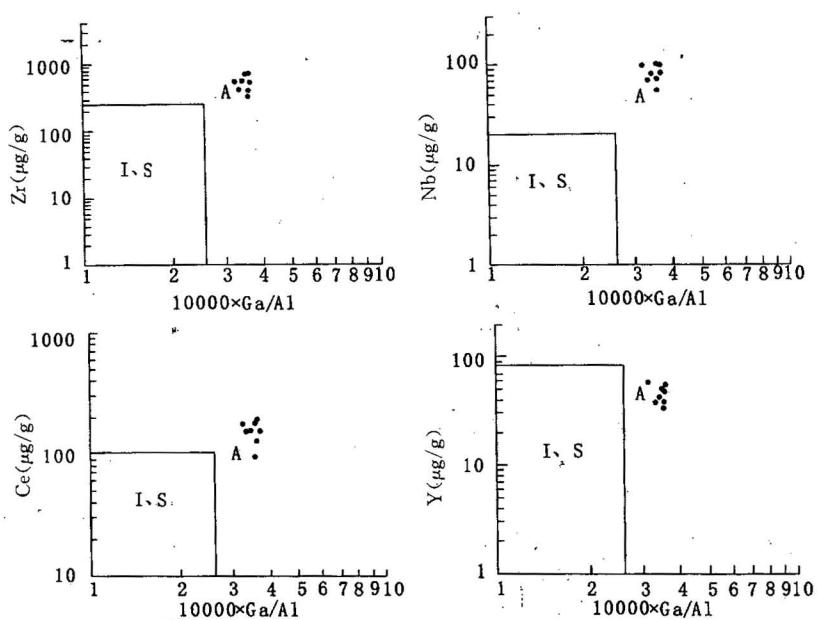


图 8 承德甲山正长岩 Zr 、 Nb 、 Ce 、 Y — $10^4 \times Ga/Al$ 关系(据 Whalen 等, 1987)
Fig. 8 Zr , Nb , Ce and Y — $10^4 \times Ga/Al$ diagrams of Jiashan syenite, Chengde
(after Whalen et al., 1987)

集, Ga 含量高, $(Ga/Al) \times 10^4$ 值变化于 3.15~3.60, 在 Whalen 等(1987)提出的多种判别图上(图 8)均落在 A 型花岗岩区。微量元素特征与瑶坑 A 型花岗岩(邱检生等, 2000)相似。

4 讨论

在 Yb — Ta 图解中(Pearce et al., 1984)(图 9),

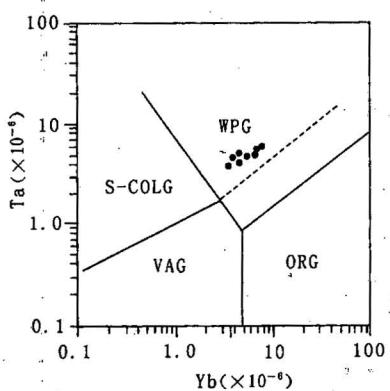


图 9 承德甲山正长岩 Yb — Ta 图解
(据 Pearce 等, 1984)

Fig. 9 Yb vs. Ta diagram for Jiashan syenite,
Chengde (after Pearce et al., 1984)

WPG—板内花岗岩; ORG—洋中脊花岗岩; VAG—火山弧
花岗岩; S-COLG—碰撞花岗岩

WPG—Within plate granites; ORG—ocean ridge granites; VAG—
volcanic arc granites; S-COLG—syn-collision granites

甲山岩体三个单元全部落在板内花岗岩区。在 SiO_2 — $Lg[Ca/(K_2O+Na_2O)]$ 图解中(Brown, 1982), 甲山岩体与燕山地区白垩纪 A 型花岗岩一样, 投影在伸展型范围, 表明岩体形成于伸展构造背景下(图 10)。R1—R2 多阳离子图解(Batchelor et al., 1985)(图 11)显示, 三个单元样品均落在造山晚期区。

地球化学特征表明, 甲山岩体属于 A 型花岗岩, A 型花岗岩可进一步划分为 A_1 型和 A_2 型, 前者为非造山 A 型花岗岩, 岩浆起源于地幔, 形成于大陆裂谷或地幔热柱、热点环境; 后者为后造山 A 型花岗岩, 岩浆起源于大陆地壳或底侵地壳, 形成于碰撞后或造山期后的张性构造环境(Eby et al., 1990, Eby, 1990, 1992)。在 A_1 、 A_2 型花岗岩判别图解(Eby et al., 1990; 图 12)中甲山岩体落在 A_1 型花岗岩区, 也落入雾灵山、千层背、山海关和响山 A 型花岗岩投点范围。洪大卫等(1995)给出的 R_1 与 Ga/Al 值图解中(图略)岩体落入 AA 区, 形成于岩石圈裂谷构造环境。从现有资料还不能断定岩体的形成是与地幔热柱或热点活动有关还是处于板内裂谷环境, 但岩浆起源于地幔, 岩体形成于伸展环境是肯定的。

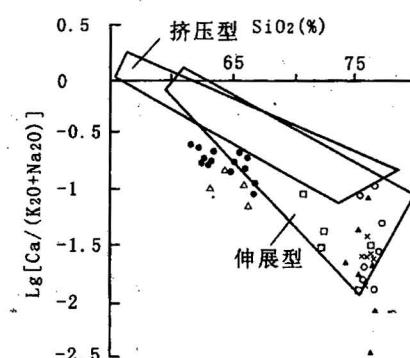


图 10 承德甲山正长岩 SiO_2 — $Lg[Ca/(K_2O+Na_2O)]$ 图解(据 Brown, 1982)

Fig. 10 SiO_2 — $Lg[Ca/(K_2O+Na_2O)]$ diagram for
Jiashan syenite, Chengde (after Brown, 1982)

●—甲山; △—雾灵山; □—千层背; ▲—山海关; ×—响山;
○—窟窿山; 除甲山外, 其它数据引自李伍平①
●—Jiashan; △—Wulingshan; □—Qianchengbei; ▲—
Shanhaiguan; ×—Xiangshan; ○—Kulongshan; data
except Jiashan are from Li Wuping①

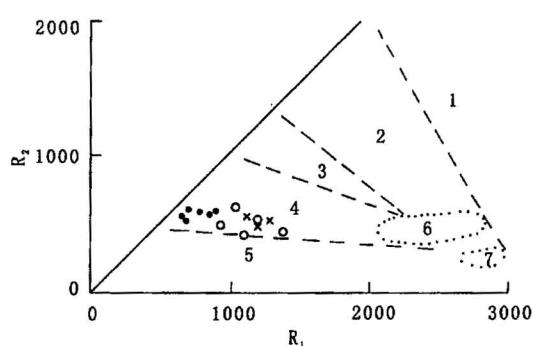


图 11 承德甲山正长岩 R_1 — R_2 多阳离子图解
(据 Batchelor 等, 1985)

Fig. 11 Diagram R_1 — R_2 multiple cations for Jiashan syenite, Chengde (after Batchelor et al., 1985)

●—姜家湾单元; ○—龙潭沟单元; ×—龙潭南沟单元; 1—幔源分异花岗岩; 2—板块碰撞前消减地区; 3—板块碰撞后隆起期; 4—造山晚期; 5—造山期后; 6—同造山期; 7—造山期后
●—Jiangjiawan unit; ○—Longtangou unit; ×—Longtannangou unit; 1—mantle fractionates granite; 2—pre-plate collision; 3—post-collision uplift; 4—late-orogenic; 5—anorogenic; 6—syn-collision; 7—post-orogenic

雾灵山正长岩经研究(李伍平^①), 其岩浆起源于岩石圈地幔, 并经过一定程度的结晶分异作用, 甲山岩体与其类似, 形成于造山晚期岩石圈伸展的拉张环境下, 岩浆活动与岩石圈减薄导致软流圈物质上涌

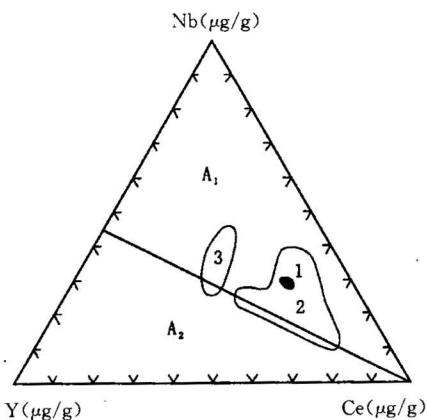


图 12 承德甲山正长岩 Nb—Y—Ce 判别图解
(据 Eby 等, 1990)

Fig. 12 Nb—Y—Ce discriminant diagram of Jiashan syenite, Chengde (after Eby et al., 1990)

A_1 —非造山的 A型花岗岩; A_2 —后造山的 A型花岗岩; 1—甲山岩体投点范围; 2—雾灵山、千层背、山海关和响山 A型花岗岩投点范围; 3—窟窿山 A型花岗岩投点范围; 2、3 据李伍平^①
 A_1 —Anorogenic A-type granite; A_2 —post-orogenic A-type granite; 1—cast dot range for Jiashan pluton; 2—cast dot range for Wulingshan, Qianchengbei, Shanaiguan and Xiangshan A-type granite; 3—cast dot range for kulongshan A-type granite; 2 and 3 after Li Wuping^①

作用有关。

早白垩世燕山区发生广泛分布的区域伸展变形(郑亚东等, 2000), 在地壳伸展环境下有八达岭地区的黑熊山和铁炉子花岗岩(王焰等, 2001), 雾灵山、千层背、山海关 A型花岗岩, 晚白垩世响山、窟窿山等碱性一过碱性的 A型花岗岩侵入(许保良等, 1994a, 1996, 1998; 李伍平^①), 揭示燕山地区白垩纪存在有地幔岩浆活动和地幔交代作用等深部地质过程(许保良等, 1994b)。

中晚侏罗世燕山地区陆内造山作用强烈, 主要为南北向地壳缩短, 导致岩石圈增厚, 大规模推覆和褶皱都发生在 161~132Ma(Davis et al., 1998)。在承德县鸡冠山下白垩统张家口组不整合覆盖在承德逆冲推覆体(Davis et al., 1998)之上, 张家口组底部熔结凝灰岩 SHRIMP 锆石 U-Pb 年龄为 135.1 ± 1Ma(赵越, 2001 未刊资料), 表明在 135Ma 前地壳挤压缩短变形作用已经结束。黄汲清等(1977)、任纪舜等(1980, 1987, 1990, 1999)、鲍亦冈等(1983)曾提出, “侏罗纪末—白垩纪初是中国东部挤压作用达到高潮, 并转向引张之开始, 是中国东部中、新生代构造发展中一个重要的转折时期”, 甲山岩体资料进一步证明了这一科学论断。从早白垩世开始, 受下地壳拆沉作用影响, 区域逐步伸展引张, 导致岩石圈减薄, 地幔物质上涌, 高钾钙碱性中基性火山岩喷发、辉绿岩脉、A型花岗岩侵入以及断陷盆地形成。伸展作用从早白垩世开始, 一直持续到晚白垩世响山和窟窿山等 A型花岗岩侵入, 晚白垩世 A型花岗岩更具伸展作用特征(图 10)。甲山岩体记录了早白垩世区域伸展和岩石圈减薄这一地质过程, 该岩体的研究对于揭示燕山地区及中国东部晚中生代岩石圈的特征和演化具有重要意义。

5 结论

(1) 承德甲山岩体岩石富碱、Si、TFe、Al、REE、Nb、Zr、Ga, 贫 Mg、Ba、Sr、Ti 和 Cr、Co、Ni 等过渡元素。Ga/Al 值大, Ce/Pb 值在 6.12~13.41, 富轻稀土, 具中等铕负异常。

(2) 岩体属 A_1 型花岗岩, 形成于造山晚期岩石圈伸展的环境下, 与地幔上涌有关。岩体位于承德晚侏罗世逆掩片东侧, 形成于早白垩世晚期, 说明区域晚侏罗世的挤压缩短和地壳增厚转变为早白垩世的伸展和岩石圈减薄, 暗示中国东部的岩石圈减薄应从早白垩世开始, 甲山岩体是燕山陆内造山带区域岩石圈减薄的早期记录。

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注 释

① 李伍平. 2002. 燕山造山带中生代火山岩地球化学特征及其地球动力学背景. 博士后研究报告, 中国科学院广州地球化学研究所.

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The Jiashan Syenite in Northern Hebei: an Early Record of Lithospheric Thinning in the Yanshan Intracontinental Orogenic Belt

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Abstract

The Early Cretaceous Jiashan syenite is located in a typical crustal thickening region, where the Late Jurassic Chengde thrust sheet was reported in the eastern segment of the Yanshan intracontinental orogenic belt. In accordance with their intruding order in time from earlier to later, the syenite can be divided into three units, the Jiangjiawan, Longtangou and Longtannangou units. Geochemically the Jiashan syenite is rich in Si, alkali, Fe, REE, Th, Ca, Nb, Zr and Hf, poor in Mg, Ba, Sr and Ti, and depleted in transitional elements such as Cr, Co, Ni, and V, with a high Ga/Al ratio. Its Ce/Pb ratio ranges from 6.12 to 13.41 and it is enriched in light REE (LREE) with a moderate Eu depletion. The Jiashan syenite, a post-orogenic intrusive, can be considered as an A-type granite, emplaced during a mantle uplifting in lithospheric extension, which suggests that the regional compressional shortening and crustal thickening ended and the lithospheric extension was prevailing. Therefore, the petrology and geochemistry of the Jiashan syenite document a lithospheric thinning in North China, which started in the Early Cretaceous.

Key words: syenite; geochemistry; tectonic implication; lithosphere thinning, Jiashan, Hebei Province