

华北侏罗纪大地构造:综评与新认识

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内容提要:论述了华北地区侏罗纪构造变形样式与岩浆活动序列,目的是试图提出一个侏罗纪构造阶段性演化历史的框架。将华北地区侏罗纪的沉积作用、构造变形和岩浆活动划分为3个阶段。早侏罗世早期大地构造显著特征是没有岩浆活动(205~191 Ma的岩浆活动空隙),华北地区处于区域性隆升,这个时期应当是华北与华南板块碰撞和陆内俯冲作用(印支运动)的延续。早中侏罗世(190~175 Ma)华北岩石圈处于弱伸展状态,表现为沿燕-辽构造带和郯庐断裂带两侧源于地幔的深成岩浆侵入、火山作用以及沿阴山—燕山构造带发育的正断层和裂谷作用,华北地块内部发生大面积整体沉降和含煤盆地的形成。这个时期大地构造仍可视为是印支运动的后续。中晚侏罗世(165±5Ma~136 Ma)构造体制发生重大变革,以多向挤压引起的强烈陆内造山和板内变形为特征。这个阶段区分出了两幕主要挤压变形事件:早幕以上侏罗统巨厚磨拉石沉积之下的角度不整合面为标志,主要沿着鄂尔多斯西缘褶皱冲断带和阴山—燕山构造带发育,时代早于160 Ma;晚幕以下白垩统底部的角度不整合面为标志,时代早于135 Ma。在中晚侏罗世构造变革时期,两个纬向构造带:北缘的阴山—燕山构造带和南缘的秦岭—大别构造带,以及鄂尔多斯盆地西缘均发生强烈的冲断褶皱变形;华北地块本身遭受NE、NNE向逆冲断层、断层相关褶皱和左旋走滑断层等构造体系的改造,并在空间分布上出现变形分解。郯庐断裂带东侧胶辽地区一系列片麻状地壳重熔型花岗岩的侵入(时代160~150 Ma)指示了中晚侏罗世重要的地壳增厚作用。华北侏罗纪变形样式、不同方向的构造体系和多向挤压是东亚周邻板块(北部西伯利亚板块、西部拉萨地块和东部古太平洋板块)从3个不同方向向亚洲大陆同时汇聚产生的远程效应的结果。正是中晚侏罗世强烈的板内挤压变形和地壳增厚作用才诱发了早白垩世(最早始于135Ma)时期岩石圈的巨量减薄和转型。

关键词: 侏罗纪构造;板内变形,华北,多向挤压;汇聚构造体系

广袤的华北地块包含不同构造地貌单元。南北两侧分别为秦岭—大别山脉和阴山—燕山山脉,西侧为狼山—贺兰山—六盘山;东侧郯庐断裂带——一条巨型岩石圈不连续构造带,纵切华北地块的东部地区。块体内部可以划分为5个构造地貌域,从西向东依次为:鄂尔多斯盆地,山西高地,北华北平原(也称大渤海湾盆地),南华北平原(也称河淮盆地),以及位于郯庐断裂东侧的胶—辽块体,各构造地貌单元之间的界线明显可辨(图1)。

从大地构造角度,华北地块属于中朝克拉通的组成部份,其结晶基底形成于(2500Ma之前)阜平运动、(2300Ma之前)五台运动和(1800Ma)吕梁运动(黄汲清,1954;Wang and Mo,1995),其克拉通台地经历了4个巨构造演化阶段:①中晚元古代沿着克拉通北缘和西缘陆内裂谷和坳拉槽的发育阶段;②古生代稳定海相台地发育阶段;③晚古生代至

早中生代期间与西伯利亚板块和华南地块相继的拼合和陆-陆碰撞阶段;④晚中生代至新生代期间克拉通岩石圈的减薄、破坏和肢解。

前人早已认识到,中生代时期华北地块发生强烈活化,经历了周期性幕式构造-热事件,导致岩石圈增厚和减薄,这种现象在山西高地以东地区尤为显著(马杏垣,1989)。然而,至今人们对巨厚而稳定的华北克拉通发生破坏和减薄的过程和机制并不十分清楚,日益成为地学研究的核心问题之一,引起了全世界地质学家的关注。侏罗纪作为东亚大地构造发展的一个重要转折时期,在中国得到了广泛的关注和研究(董树文等,2000)。翁文灏(Wong, 1927,1929)最早提出的“燕山运动”,专指侏罗纪构造变形事件,已经成为地质学界最广泛引用的构造术语。目前对这一概念的内涵得到了扩大和提升,它包含了侏罗纪至白垩纪之间发生的所有构造-热

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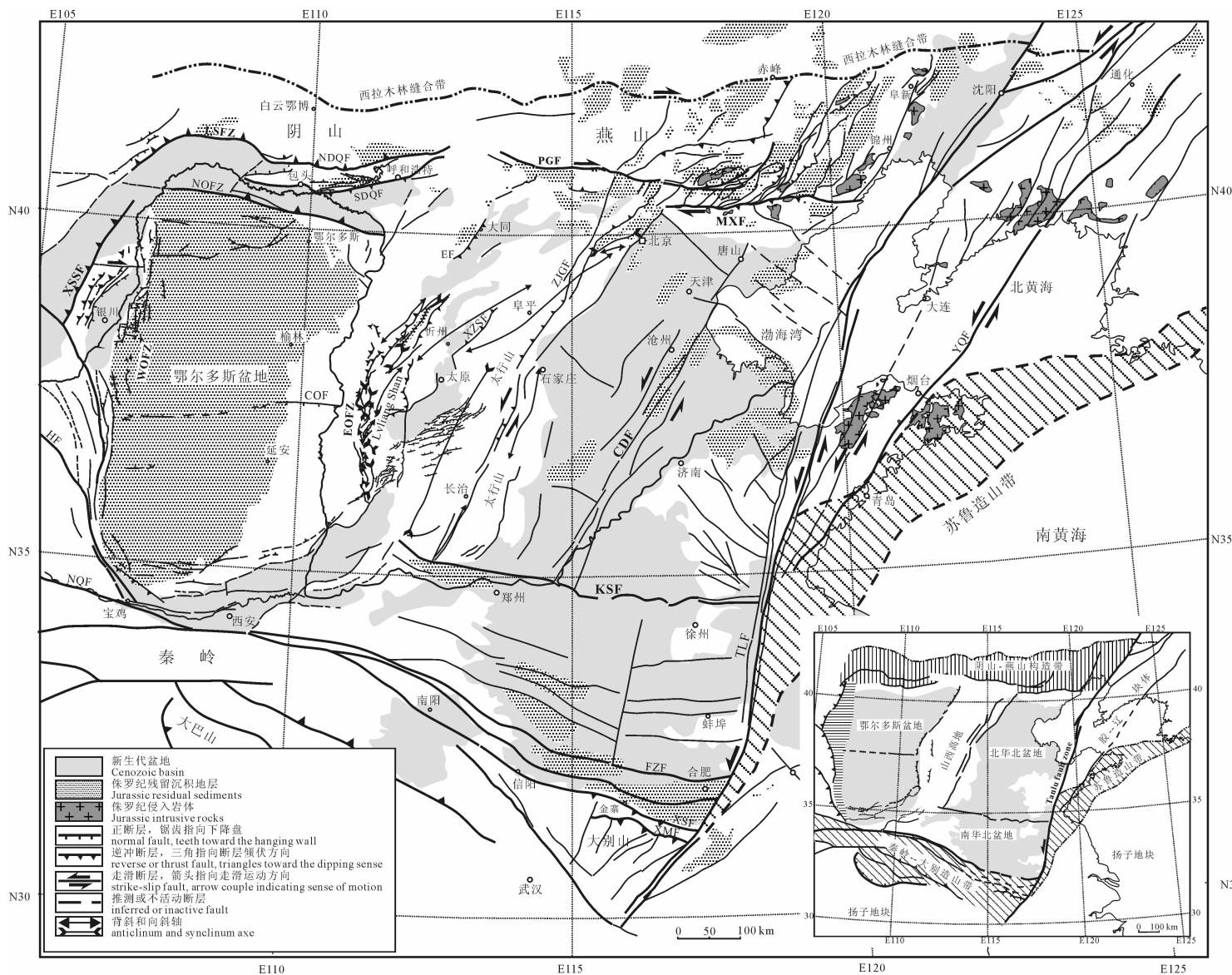


图 1 华北侏罗纪构造纲要图

Fig. 1 Outline of Jurassic tectonics of North China

右下角插图: 华北主要构造地貌单元 Inset at lower right corner presents main physiographic units of North China

TLF—郯庐断裂带; YQF—鸭绿江—青岛断裂带; PGF—平泉—承德断裂; MXF—密云—喜峰口断裂带; ZJGF—紫金山断裂带; CDF—沧东断裂; EF—鹅毛口断裂; EOFZ—鄂尔多斯东缘断裂带; XZSF—系舟山断裂; COD—鄂尔多斯中央断裂; SDQF—南大青山断裂带; NDQF—北大青山断裂带; NOFZ—鄂尔多斯北缘断裂; LSFZ—狼山断裂带; XSSF—小松山断裂; KSF—开封—商丘断裂; HF—海原断裂; NQF—秦岭北缘断裂; FZF—肥中断裂; XSF—信阳—舒城断裂带; XMF—晓天—磨子潭断裂

TLF—Tan-Lu fault zone; YQF—Yalvjiang-Qingdao fault zone; PGF—Pingquan-Chengde fault zone; MXF—Miyun-Xifengkou-Jingxi fault zone; ZJGF—Zijinguan fault zone; CDF—Cangdong fault zone; EF—Emaokou fault; EOFZ—Eastern Ordos fault zone; XZSF—Xizhoushan fault zone; COD—Central Ordos fault; SDQF—Southern Daqingshan fault zone; NDQF—Northern Daqingshan fault zone; NOFZ—Northern Ordos fault zone; LSFZ—Langshan fault zone; XSSF—Xiaosongshan fault zone; KSF—Kaifeng-Shangqiu fault zone; HF—Haiyuan fault zone; NQF—northern Qinling fault zone; FZF—Feizhong fault; XSF—Xinyang-Shucheng fault zone; XMF—Xiaotian-Mozitan fault zone

事件(黄汲清, 1954; 任纪舜等, 1990; 崔盛芹等, 2002; 董树文等, 2007)。

本文收集了近 20 年来不同学科(沉积学、石油地质学、构造地质学、岩石学、地质年代学等)获得的研究成果和重要的岩体年代学测试数据, 在整理分析这些文献和数据资料的基础上, 结合作者近年来的研究成果, 编制了华北侏罗纪构造纲要图, 重点剖析了华北地块侏罗纪地层、构造变形和岩浆作用, 旨在从一个区域宏观的角度阐述华北地块侏罗纪板内变形和岩浆作用过程、构造体制转换及其板块和深部地球动力学背景。

1 侏罗纪地层概述

侏罗纪地层主要保存在山西高地以西的鄂尔多斯盆地、南华北南部的合肥盆地、沿着阴山—燕山构造带发育的地堑盆地中。在北华北平原之下也钻遇了侏罗纪地层, 但其年代学格架研究程度较低(漆家福等, 2003, 2004)。表 1 汇总了基于地层接触关系、层序分析和同位素年代资料汇总的华北地块地层单元对比。燕山—辽西构造带和北京西山的侏罗纪地层中发育大量的火山岩。前人对这些火山岩进行了大量的 ^{40}Ar — ^{39}Ar 和 SHRIMP 年龄测试, 这些测试数据成为建立华北地区侏罗纪地层年代格架的基础。总体来说, 侏罗纪地层包括上、中、下侏罗统三个序列, 地层之间或为侵蚀接触, 或为角度不整合接触, 它们构成了划分华北地区侏罗纪阶段性构造演化历史的基础。

1.1 下侏罗统

在鄂尔多斯盆地内部, 下侏罗统包括富县组和延安组 2 个地层单元。最下部的侏罗系随着华北板块和华南地块三叠纪碰撞引起的隆升剥蚀而缺失(刘池阳等, 2005, 2006)。富县组对应于一套河流、湖泊相的沉积, 延安组为一套厚度稳定且富煤的沉积地层, 总厚度在 250~400 m。这套地层可与山西北部大同和宁武残留盆地中的侏罗系对比(程守田等, 1997)。古地理重建结果显示, 下侏罗统地层和富煤层序覆盖了整个鄂尔多斯盆地和山西高地, 向东或许延伸到华北平原之下的广泛范围内(王双明等, 1999)。

大青山地区发育的下侏罗统以五当沟组为代表, 沉积在受南缘北倾主干正断层控制的引张盆地中(Ritts et al., 2001)。200~300m 厚的砾岩沿边界断裂上盘堆积, 向上被河流—湖泊相和三角洲相的砂岩、粉砂岩、泥岩所覆盖, 形成一个递进沉积序

列(彭向东等, 2001)。

在合肥盆地, 沿着盆地南部边缘出露的下侏罗统命名为防虎山群, 对应一套粗粒砂岩、粉砂岩和泥岩沉积, 厚度大约为 1000m(李双应等, 2002, 2005)。这套地层在合肥盆地内被安参—1 井钻遇, 厚约 230 m, 主要由泥岩和泥质粉砂岩组成(徐佑德等, 2002), 代表半深水的湖泊沉积环境。

在北京西山和沿燕—辽构造带, 下侏罗统以火山岩为主, 在北京西山和燕山西段为南大岭组, 在辽西地区以兴隆沟组或水泉沟组为代表。火山喷发时代在 180~194 Ma(赵越等, 2002, 2006; 徐刚等, 2005)。火山岩系被北京西山窑坡组和辽西北票组等含煤碎屑沉积所覆盖, 这些地层可与鄂尔多斯盆地中的下侏罗统延安组对比。

1.2 中侏罗统

中侏罗世的沉积环境发生了改变, 在鄂尔多斯盆地以湖水逐渐变深的湖泊相沉积为特征(程守田等, 1997; 刘池阳等, 2006)。包含两个组: 下部直罗组由砂岩、粉砂岩和泥灰岩组成的湖泊相沉积地层, 不含煤层, 厚约 200~300 m, 基底以指示早中侏罗世盆地上隆的剥蚀界面为标志; 上部安定组由泥岩夹层、钙质粉砂岩、泥灰岩和白云岩等组成, 厚约 50~400 m, 指示干燥的湖泊沉积环境。

大青山盆地的中侏罗统地层以长汉沟组为代表, 包含由浅水到深水的湖泊相沉积, 由灰色到绿色的粉砂岩、砂岩、钙质粉砂岩、泥灰岩、碳酸盐岩以及凝灰质砂岩组成。

在合肥盆地, 中侏罗统地层在安参-1 井钻探中显示与下侏罗统为连续沉积, 总厚度达 2407 m(徐佑德等, 2002), 主要由泥岩和砂岩互层组成。

在北京西山和沿燕—辽构造带, 中侏罗统由一套递进沉积序列组成, 下部为砾岩与粗砂岩, 向上变为砂岩到粉砂岩, 可与鄂尔多斯盆地中侏罗统地层对比。在北京西山, 龙门组底部为剥蚀间断面, 九龙山组底部为不整合面, 指示中侏罗世晚期该区快速抬升(赵越等, 2002)。

1.3 上侏罗统

鄂尔多斯盆地上侏罗统局限在盆地西缘, 以芬芳河组为代表, 对应一套巨厚的砾岩沉积(张抗, 1989, 见表 1)。这一地层出露于鄂尔多斯盆地西缘南部的芬芳河谷中, 地震剖面上表现为楔形体, 南厚、北薄(刘池阳等, 2006), 以角度不整合超覆于早中侏罗世和其他较老的地层之上, 并被下白垩统砾岩层所超覆。砾石成分混杂, 以变质岩、脉石英、

表 1 华北地区侏罗纪地层组对比

Table 1 A correlation of Jurassic Formations in North China

地 层				鄂尔多斯盆地	大青山 (石拐盆地)	山西北部		燕山-辽西构造带			合肥盆地	
年龄 (Ma)	纪	世	阶	W	E	大同西部	北京西山	冀北	辽西	内蒙古		
130	Cretaceous	L. Cretaceous		志丹群		左云组		张家口组 ▽▽▽▽▽ (~135 Ma)			周公山组 ▽▽▽▽▽	
			Hauterivian									
140			Vallanginian									
			Berriasian	芬芳河组 (砾岩层)	大青山组 (砾岩层)		后城组 (砾岩层) (149~137Ma)	后城组 (砾岩层) (152~136 Ma)	土城子组 (砾岩层) (139 Ma)	土城子组 (砾岩层)	珠肌组 (砾岩层)	
150	U. Jurassic	Tithonian				髫髻山组 ▽▽▽▽▽ (157Ma)	髫髻山组 ▽▽▽▽▽ (153~158 Ma)	兰旗组 ▽▽▽▽▽ (158~154 Ma)	髫髻山组 ▽▽▽▽▽ (152 Ma)			
		Kimmeridgian										
160		Oxfordian										
	M. Jurassic	Callovian		安定组 直罗组	长汉沟组	云岗组	龙门组		Deformation event-I	北票组	圆桶山组	
		Bathonian										
170		Bajocian										
		Aalenian										
180	L. Jurassic	Toarcian	延安组 (含煤)	五当沟组 第4段 第3段(含煤层) 第2段(含煤层) 第1段	大同组 (含煤层)	永定庄组	南大岭组 ▽▽▽▽▽ (174~190 Ma)	南大岭组 ▽▽▽▽▽ (180 Ma)	兴隆沟组 ▽▽▽▽▽ (188~194 Ma)		防虎山组	
		Pliensbachian	富县组									
190		Sinemunan										
		Hettangian					区域隆升与剥蚀					
200	Triassic	U. Triassic	Rhaetian									
			Norian									延长组
210												
根据 Gradstein et al., 2004				据程守田等, 1997, 刘池阳等, 2006	据彭向东等, 2001	据程守田等, 1997	赵越等, 2002, 2006 Yuan et al., 2005	Cope, 2003; 刘建等, 2006	Swisher et al., 2005 赵越等, 2002 Yang et al., 2005	柳永清等, 2004	李双应等, 2002; 2005	



火山岩层



同构造沉积地层

碳酸盐岩、砂岩等为主,它代表了一套伴随鄂尔多斯西缘逆冲褶皱构造带发育的同构造生长地层。

在大青山的石拐盆地,上侏罗统称为大青山组,对应一套较厚的砾岩沉积,底部以角度不整合超覆在中晚侏罗统之上(图 2b),总厚度为 332 m(彭向东等,2003)。这套砾岩层代表了同构造生长地层,伴随着中晚侏罗世大青山褶皱冲断带的发育而堆积。

沿燕-辽构造带,上侏罗统发育在火山盆地之中,可以划分为上下两个组:下部的髫髻山组或蓝旗组和上部以砾岩为主的后城组或土城子组。前人对这两个地层组的年龄做了大量的工作,Davis(2005)、刘建等(2006)、孙立新等(2007)做了相继的述评。髫髻山组或蓝旗组中的火山岩多为中酸性岩石组合,其中夹有碎屑岩层,在不同盆地所做的火山岩⁴⁰Ar-³⁹Ar 和 SHRIMP 年龄测试结果集中在 153~165 Ma(Cope, 2003; Swisher et al., 2001; 柳永清等,2004; Gao et al., 2004; 张宏等,2002; 徐刚等,2005; 刘建等,2006;孙立新等,2007)。较厚的后城组或土城子组沉积在气候干燥的河流—湖泊相环境之下,包含 3 个岩性组合:下部为紫色粉砂岩、页岩、砂岩和碳酸盐岩互层;中部为凝灰质泥岩、砾岩;上部为砾岩和砂岩,沉积厚度在 1000~3000 m(和政军等,1998)。其中夹有火山岩层,测年结果在 152~136 Ma(刘建等,2006)。

2 侏罗纪变形样式和构造体系

华北地块侏罗纪变形非常广泛,前人对此进行了大量的调查和研究。侏罗纪变形是不均匀分布,在空间上发生变形分解。图 1 展示了华北地区典型的侏罗纪构造,这些包括鄂尔多斯盆地西缘的褶皱冲断带(WOTFB)、山西逆冲褶皱系统(STFS),郯庐走滑断裂系(TSFS),两个纬向构造带:北侧的阴山—燕山构造带和南侧的秦岭—大别构造带,均受逆冲推覆作用的改造活化。

2.1 鄂尔多斯盆地西缘褶皱冲断带(WOTFB)

该带沿鄂尔多斯西缘发育,在东经 106° 和 107° 之间沿 NS 向延伸,NS 向长 600km,EW 向宽 30~80 km。该构造带被视作为亚洲中部典型的板内变形带(Darby et al.,2002)。众多学者从地球物理、钻孔资料(杨俊杰,2002; 赵重远等,1990)以及野外调查的角度(张抗,1989; 张岳桥等,2006,2007)对该区进行了详细的研究。从展布形态看,这一“Z”字形展布的构造带可以划分为 3 段,从北向南依次为:桌子山和横山堡段,中间马家滩段和南段,3 段

之间构造差异显著。桌子山复合冲断带卷入前寒武纪结晶基底,典型构造样式表现为向东逆冲的叠瓦状断裂组合,如桌子山断裂,主干断裂向西陡倾,倾角在 60°~70°,断距从千米以上至数百米不等,缩短率达 32%。南段叠瓦冲断带以盖层的拆离和滑脱为特征,其根带位于青铜峡—固原断裂带和海原—六盘山断裂带,在马家滩地区,缩短量达 33.6%~42.8%。相对而言,中段冲断带的构造样式较特殊,它由一组向东倾的阶状逆冲断层组成,断距大者超过千米,但缩短率仅为 10%。由于受到新生代银川地堑发育的影响和后期的改造作用,该阶状冲断带与贺兰山褶皱造山带之间的构造关系并不十分清楚,因而出现不同的构造解释。大多数学者认为,贺兰山北端的 WE 向正宜关断裂向东延伸与桌子山冲断带南端一条横向断裂相连,该断裂与贺兰山褶皱系和西缘冲断带发育过程中起着重要的横向位移调节作用。但根据 Darby 等(2002)的解释,正宜关断裂与桌子山北端的一条横向断裂原是同一条,后被近 NS 走向的黄河断裂左旋错移了近 60 km。这些作者根据这种推断进行的位移复原图得出,桌子山冲断带是贺兰山褶皱带的前陆部分。

根据断层发育特征与接触关系,以及褶皱逆冲造带卷入的最新地层推断,西缘冲断带主要形成于中侏罗世晚期至晚侏罗世时期的燕山运动。沿盆地西缘发育的晚侏罗世芬芳河组巨厚砾岩层(1200~3000 m),代表了同构造期的沉积记录;砾岩层向东陡倾,其上被早白垩世砂砾层不整合超覆。贺兰山地区侏罗纪地层河流流向和物源分析表明,贺兰山盆地发生倒转的时间在中侏罗世沉积之后。另外,野外调查显示,贺兰山褶皱冲断带西缘的小松山断裂,奥陶纪碳酸盐岩逆掩到中侏罗世直罗组和安定组之上(图 2c),表明逆冲推覆作用发生在中侏罗世沉积之后。

2.2 山西台褶带(STFS)

该带位于华北地块中部,构造形迹包括 NNE 至 NE 走向的逆冲断层,以及山西高地“S”型展布的断层相关褶皱。沿着高地轴向发育的晚新生代山西地堑系,其多数边界正断层是由中晚侏罗世逆冲断裂转变而成。以系周山断层为例(图 1 中 XZSF),断层边界位于忻县地堑东缘。据孙战良等(2004)的野外调查,这是一条晚中生代逆冲断裂,逆冲方向由西向东,估算最小水平位移量为 5.8 km(图 2i)。逆冲断层倾向东或西,卷入了前寒武纪变质基底至中侏罗世沉积盖层。断层破碎带和角砾指示以脆性

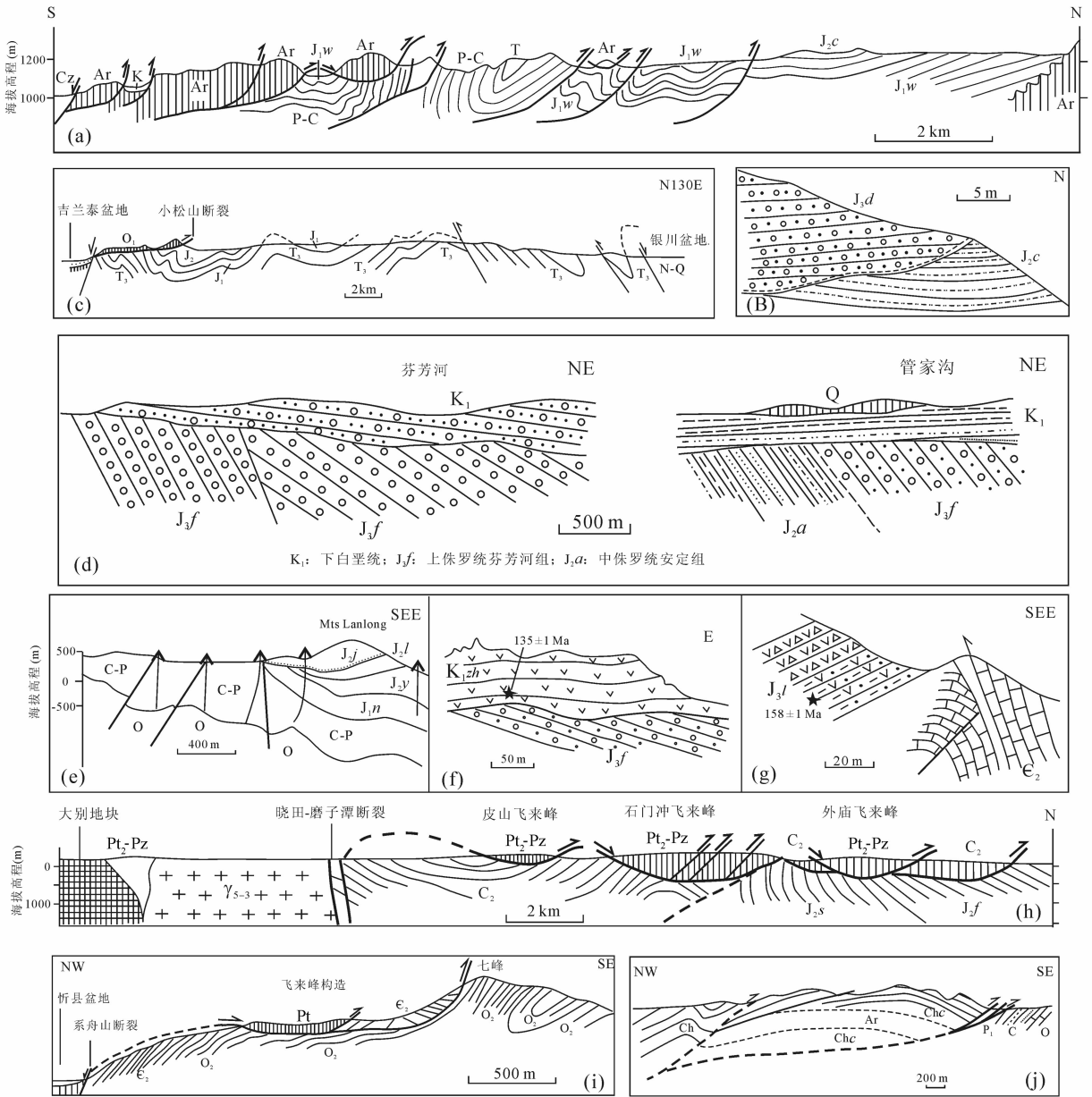


图 2 华北地区侏罗纪变形事件与构造样式简图

Fig. 2 Sketches showing Jurassic deformation events and structural styles in North China

(a)一大青山南北向构造横剖面简图(据杜菊民等,2006);Cz—新生界;K—白垩系;J_{1w}—下侏罗统五当沟组;J_{2c}—中侏罗统长汉沟组;P-C—石炭—二叠系;Ar—太古宇;(b)—上侏罗统大青山组和中侏罗统长汉沟组之间的不整合接触(据彭向东等,2003);(c)—简化贺兰山 WE 向构造剖面图(张抗,1989);N-Q—新近系—第四系;J₂—中侏罗统;J₁—下侏罗统;T₃—上三叠统;O₁—下奥陶统;(d)—鄂尔多斯西缘上侏罗统芬芳河组底部不整合面(根据张抗,1989);(e)—京西中侏罗统九龙山组之下的不整合面(据赵越等,2003,2004);(f)—位于早白垩世张家口组底部的角度不整合面;(g)—辽西陵源盆地晚侏罗世蓝旗组之下的不整合面(据赵越等,2003,2004);(h)—北大别北淮阳构造带南北向构造横剖面简图(根据刘文灿等,1999);J_{2s}—中侏罗统三尖铺组;J_{2f}—中侏罗统凤凰岭组;C₂—中石炭统;Pt₂-Pz—元古宙至古生界;(i)—山西北部系舟山断裂带构造剖面(据孙战良等,2004);(j)—北京西山南大寨逆冲推覆构造剖面图(据赵占元等,1994)

(a)—Simplified N-S structural section across the Daqingshan belt (modified from Du et al. , 2006); Cz—Cenozoic; K—Cretaceous; J_{1w}—Lower Jurassic Wudanggou fm.; J_{2c}—Middle Jurassic Changhangou fm.; P-C—Permo-Carboniferous; Ar—Archean; (b)—Sketch of the disconformity between Upper Jurassic Daqingshan Formation and Middle Jurassic Changhangou Formation (after Peng et al. , 2003); (c)—Simplified structural section across the E-W Helanshan (after Zhang, 1989); N-Q—Neogene-Quaternary; J₂—Middle Jurassic; J₁—Lower Jurassic; T₃—Upper Triassic; O₁—Lower Ordovician; (d)—Unconformity beneath the Upper Jurassic Fengfanghe formation along the western margin of the Ordos basin (after Zhang Kang, 1989); (e)—Stratigraphic unconformity beneath the Middle Jurassic Jiulongshan formation in West Beijing; (f)—Stratigraphic unconformity beneath the Early Cretaceous Zhangjiakou Fm. in west Yanshan; (g)—Stratigraphic unconformity beneath the Upper Jurassic Lanqi Fm. in Lingyuan basin, Liaoxi (after Zhao Yue et al. , 2003, 2004); h—Simplified structural section across N-S northern Huaiyang belt, north of the Dabie belt (after Liu et al. , 1999); J_{2s}—Middle Jurassic Sanjianpu fm.; J_{2f}—Middle Jurassic Fenghuangling fm.; C₂—Middle carboniferous; Pt₂-Pz—Proterozoic to paleozoic; i—structural section of the Xizhoushan fault, north Shanxi (after Sun et al. , 2004); j—structural section of the Nandazhai thrust nappe in West Beijing (after Zhao and Shan, 1994)

变形为主。褶皱构造多呈箱状,受前寒武纪基底中的滑脱构造所控制。早中侏罗世地层保存在残留的向斜盆地中,如山西地堑系西缘的大同盆地和静乐盆地、太行山南缘的沁水盆地等。小规模早白垩世岩体侵入其中,但未受挤压变形的影响。廖昌珍等(2007)和张岳桥等(2007)基于对鄂尔多斯盆地东缘和山西断褶带西部地区构造样式的分析结果,认为整个山西台褶带的构造样式可以解释为与山西高地深部向西扩展的基底滑脱构造有关。详细的遥感图像和野外调查表明,这一复杂地貌边界带表现为一系列向西陡倾的冲断层和上盘断坡,指示鄂尔多斯东缘构造带受控于深部背冲式的滑脱构造。

2.3 郯庐走滑断裂系

郯庐走滑断裂系,也称为郯庐扭动断裂系(Xu et al., 1987),是指 NNE-SSW 走向、近平行于郯庐断裂带、不同规模断层或断裂带组合的总称,它纵切了东亚大陆东部,构成了东亚大陆东部特征的陆缘走滑断裂系。除了错移大别—苏鲁超高压变质带的郯庐断裂带以外,位于郯庐断裂东侧、切穿胶辽块体的鸭绿江—青岛断裂(YQF,图 1)也是一条重要的地壳尺度断裂带。郯庐断裂带本身在三叠纪华北与华南地块碰撞阶段,表现为走滑特征(Zhu et al., 2005;朱光等,2006)。郯庐断裂系的一个共同特征是其左旋走滑运动性质,并认为左旋走滑主要发生在晚三叠世至早侏罗世时期,尽管这个时期走滑位移量的估算一直是一个未曾解决的问题。由于白垩纪时期广泛的伸展和沉降(朱光等,2002;Zhang et al., 2003),郯庐断裂带侏罗纪变形特征很难直接观察和研究。然而,沿大别山东缘、郯庐断裂带东缘的张八岭—肥东块体和苏鲁带的西缘,陆续发现了左旋韧性剪切变形形迹。通过韧性剪切带中角闪石、多硅白云母和黑云母的 $^{40}\text{Ar}/^{39}\text{Ar}$ 测试,获得的年龄为 136~143Ma(Zhu et al., 2005)和 162±1Ma~156±2Ma(Wang Y., 2006),这一年龄代表了剪切带在晚侏罗世不同时间段的冷却时代。另外,在胶辽块体中广泛发育的晚侏罗世“S”型花岗岩体,结晶年龄在 165~150 Ma(苗来成等,1998;关康等,1998;郭敬辉等,2005),其特征的片麻理构造指示了晚侏罗世花岗岩体侵位过程的韧性变形。在胶北隆起区,晚侏罗世花岗岩体侵位后期的脆性变形十分明显,表现为一系列 NW 倾向的低角度逆冲断层,时代早于早白垩世伸展之前,限定在 150~135Ma 之间。杨进辉等(2004)利用激光 $^{40}\text{Ar}/^{39}\text{Ar}$ 测年技术,对辽东半岛晚侏罗世侵入岩的挤压变形

时代进行了测试,获得的结果显示岩体韧性变形发生在约 143 Ma,早于早白垩世伸展作用(128 Ma)。叶慧文等(1994)报道了一组脉状青铝闪石 $^{40}\text{Ar}/^{39}\text{Ar}$ 年龄测试结果,约 154Ma,样品来自于东北牡丹江地区沿敦化—密山断裂带出露的蓝闪石片岩之中。所有这些研究成果和测年结果表明晚侏罗世郯庐断裂带以左旋走滑剪切变形为主。

2.4 阴山—燕山构造带

阴山—燕山构造带位于华北地块的北缘,经历了多阶段的演化历史。北部边界的西拉木伦(或索伦)缝合带将华北板块和兴蒙造山带相分割。沿阴山—燕山构造带,侏罗纪挤压变形形迹十分显著。

燕山构造带是中国中生代构造研究最为详细的地区。基于该区地层角度不整合接触关系,翁文灏(Wong, 1927, 1929)将“燕山运动”的主要变形事件划分为“A”幕、“B”幕和“中间”幕。根据侏罗纪火山岩的同位素年代学研究(赵越等, 2002, 2004;徐刚等, 2005;刘建等, 2006),将“A”幕(髫髻山火山岩下部的角度不整合)作为该时期构造变形的最强烈阶段,起始时代为 160±5 Ma;“B”幕以早白垩世张家口组火山岩底部的角度不整合为代表、以向南倾的逆冲作用为特征,时代早于 135.8±3.1Ma(牛宝贵等,2003;Niu et al., 2004);“中间”幕表现为髫髻山组或蓝旗组,时代为 160~139Ma。

Davis 等(1996, 1998, 2001)和郑亚东等(2000)对燕山构造带进行了系统的研究,并识别出几个变形幕。最早变形幕(phase-I 变形)记录在平泉关—古北口断裂带向南倾的逆冲断层(PGF,图 1),时代早于 180Ma,很可能界于晚三叠纪至早侏罗世之间。燕山构造带最主要的挤压变形发生在晚侏罗世时期(phase-III 变形)。该变形时期形成的许多逆冲推覆构造被识别和鉴定出来,包括向北至北西向逆冲的承德推覆体,时代为 161~148 Ma,十三陵逆冲带时代为 161~141 Ma,云蒙山褶皱和变质时代为 141~143 Ma,向南倾的四合堂和古北口逆冲推覆构造时代起始于 148~143 Ma,持续到 127~132 Ma。

通过对沿阴山—燕山构造带晚侏罗世沉积特征研究,和政军等(1998, 1999)推断,沿阴山—燕山前陆带发育一个延伸 1200km 长的盆地群,并将其与阴山—燕山构造带向南逆冲推覆作用联系起来。这一假设受到了 Davis(2005)的质疑,他认为晚侏罗世挤压作用既有向南逆冲,也有向北逆冲。基于对燕山构造带中段和辽西地区的逆冲推覆构造和古

北口—平泉断裂带(PGF, 图 1)和密云—喜峰口断裂带(MXF, 图 1)两个主要 WE 走向的断裂带几何学和内部结构的分析,张长厚等(2001,2002)建立了一个统一的运动学模型来解释该区复杂的构造变形特征。这些作者研究发现,近 WE 走向的主干断裂带以右旋走滑的方式横切燕山构造带中部,向东延伸至辽西地区,与北东向的逆冲断层及其相关褶皱相连,它们共同形成了一个双重右旋走滑系统,其主要活动时期在 148~132Ma。这一运动学模型与内蒙古赤峰地区的西拉木伦(索伦)右旋走滑活动相一致,刘伟等(2003)报道了在这一地区西拉木伦(索伦)右旋走滑变形的一个 Rb-Sr 同位素测年结果为 165Ma,这一年龄代表了该断裂带韧性变形的时代。

沿大青山构造带发现了大规模的逆冲岩席,主要发育在大青山北缘断裂和南缘断裂带之间的狭长地带(NDQF, SDQSF, 图 1)。大青山北缘断裂带作为内蒙古地轴的南界,被认为向南逆冲,而大青山南缘断裂带向北逆冲,发生的时代早于早白垩世伸展构造。在这两条断裂带之间地带,发育一系列侏罗纪断陷盆地和大型逆冲推覆构造,陈志勇等(2002)和刘正宏等(2002)对其进行了详细的调查和研究。根据郑亚东(Zheng et al., 1998)对大青山东部逆冲岩片的几何学和运动学研究结果,该区前寒武纪变质基底岩石逆掩到早中侏罗世沉积地层之上,形成飞来峰和构造窗构造,估算最小位移量为 25~30 km,根带位于大青山南缘断裂带。在大青山西部地区,将北部边缘的逆冲系统划分位 3 个构造区带(杜菊民等,2005),从南向北依次为:逆冲叠瓦带、同斜至倒转的褶皱逆冲带、断层相关褶皱带,变形强度由南向北依次减弱。逆冲方向由 SSE 向 NNW,最小位移量为 10~20 km。由于早中侏罗世地层卷入到逆冲推覆带中,由此推断构造运动时代为晚侏罗世。阴山构造带最南部边界位于鄂尔多斯盆地北缘断裂带,这是一条向北倾的正断层(NOFZ, 图 1),构成了新生代河套地堑盆地的南界。刘正宏等(2004)沿着鄂尔多斯北缘详细的野外调查,发现一个小规模的前寒武纪变质岩片推覆到二叠系之上,估算最小水平位移量为 5km。这进一步说明,隐伏在河套地堑盆地之下的基底曾是一个 W-E 向延伸的构造凸起或高地,在晚侏罗世受到南北两侧边界断层的反向逆冲所致。

2.5 秦岭—大别构造带

该构造带演化历史相当复杂,主要表现为沿华北地块南缘的多阶段碰撞和增生历史(张国伟等,

1996, 1997; Ratschbacher et al., 2003)。三叠纪时期华北与华南地块的拼合碰撞和陆内深俯冲作用形成了这一巨大的纬向构造带,同时沿大别—苏鲁构造带露出了世界上分布面积最大的超高压变质岩。侏罗纪变形虽然重要,但研究程度较低。最好的侏罗纪变形在大别山北部的北淮阳构造带有发现,沿该带不同的变质岩系和侏罗纪岩层被推覆到晚侏罗世沉积序列之上,并被早白垩纪火山岩所覆盖(孙晓孟等,2004a,b)。最为强烈的缩短变形发生在晓天一磨子潭断裂带(XMF)和信阳—舒城断裂带(XSF)之间(刘文灿等,1999)。主干逆冲断层的倾向指示了 NNE-SSW 向的挤压。

3 侏罗纪岩浆作用

侏罗纪侵入岩体主要出露于郯庐断裂以东的胶莱—辽东块体、郯庐断裂带以西的徐-淮地区以及燕山—辽西地区。对这些地区侵入岩体高精度年代学测试数据资料进行了收集和整理,从而建立了中生代岩浆作用的年代序列,见散点图 3 和图 4。

出露于胶辽地区和蚌埠隆起带的大多数片麻状花岗岩,以往的区调认为该类岩体形成于元古宙,而通过高精度 SHRIMP 测年手段的应用,确定了这些岩体的侵入时代为 150~160 Ma,即晚侏罗世。基于胶辽和徐淮地区高可信度岩体年龄测试数据分析(图 3),表明侏罗纪岩浆序列与晚三叠世和早白垩世岩浆序列之间的界限是明显的。在晚三叠世与早侏罗世早期之间,存在一个明显的岩浆活动间隙期,时代在 205~190 Ma,它很可能对应着三叠纪时期华北与华南地块碰撞和陆内深俯冲的最后阶段。晚侏罗世岩浆序列与早白垩世岩浆序列之间似乎也存在一个很短暂的岩浆活动间隙期,时间界于 143~136 Ma。总体来说,胶辽地区和蚌埠隆起带侏罗纪岩浆活动包含两个岩石组合:早中侏罗世辉长岩和闪长岩,晚侏罗世“S”型花岗岩。早侏罗世侵入岩体年龄位于 190~175Ma,岩石类型和地球化学特征指示岩体形成于伸展构造环境下幔源岩浆的上升侵位。晚侏罗世岩体形成时代集中在 160~150 Ma,个别形成于 172~165 Ma 和 149~142 Ma 之间。岩石类型、地球化学特征和野外观察表明岩体属于壳源岩浆作用,明显不同于早白垩世壳幔混合型侵入岩体。

沿燕山—辽西构造带,侏罗纪岩体主要分布在辽西地区。岩体年龄数据的散点图分析显示(图 4),在 210~195 Ma 存在一个岩浆活动的间隙时

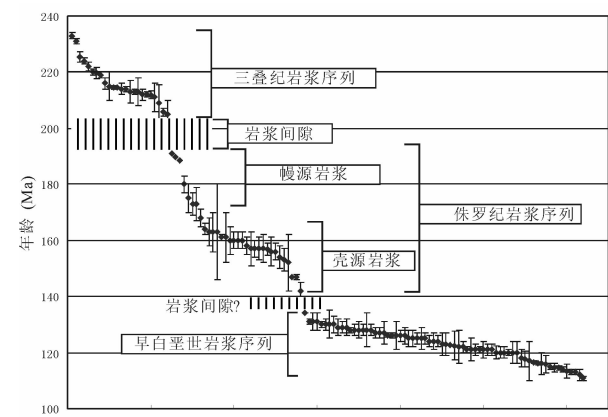


图 3 胶—辽地区与鲁西—徐淮地区中生代侵入岩同位素年代学数据散点图

Fig. 3 Scatter plot of isotopic dating ages of Mesozoic intrusive rocks from the Jiao-Liao Massif and Luxi-Xuhuai regions

数据来源:Yang et al. (2004, 2005); Chen et al. (2003); 郭敬辉等(2005); 苗来成等(1998); 胡世玲等(1987); 胡华斌等(2004); 徐洪林等(1997); 关康等(1998); 赵广涛等(1997); Zhao et al. (1998); 周建波等(2003); 吴福元等(2005, 2006); Wu et al. (2005); 林景任等(1996, 2000); 靳克等(2003); 许文良等(2004); Xu et al. (2005); 杨承海等(2005); 杨德彬等(2005)

Data sources from: Yang et al. (2004, 2005); Chen et al. (2003); Guo et al. (2005); Miao et al. (1998); Hu et al. (1987); Hu H. et al. (2004); Xu H et al. (1997); Guan et al. (1998); Zhao et al. (1997, 1998); Zhou et al. (2003); Wu et al. (2005, 2006); Lin et al. (1996, 2000); Jin et al. (2003); Xu W et al. (2004, 2005); Yang C. et al. (2005); Yang D. et al. (2005)

期。侏罗纪岩浆序列和早白垩世岩浆序列之间的界线不是很清楚。总体来说,这个地区侏罗纪岩体可以划分为三个时间阶段:早侏罗世(195~180 Ma)以辉长岩-闪长岩-石英闪长岩-花岗岩组合为特征,中侏罗世(180~170 Ma)以闪长岩-石英闪长岩-花岗岩组合为特征,而晚侏罗世(165~138 Ma)则以花岗岩为主。在 170Ma 和 165Ma 之间似乎存在一个短暂的岩浆活动间隙期。在燕山西部地区,侏罗纪岩体出露较差,仅有部分岩体出露于云蒙山地区,测定的侵入时代在 160~141Ma (Davis et al. , 2001,2003)。这些作者认识到,这地区包含两个岩石组合:一个为起源于原始地幔的辉长岩和闪长岩,另一个为壳源的花岗闪长岩。

4 华北侏罗纪大地构造演化阶段

以上的介绍和分析表明,华北侏罗纪大地构造

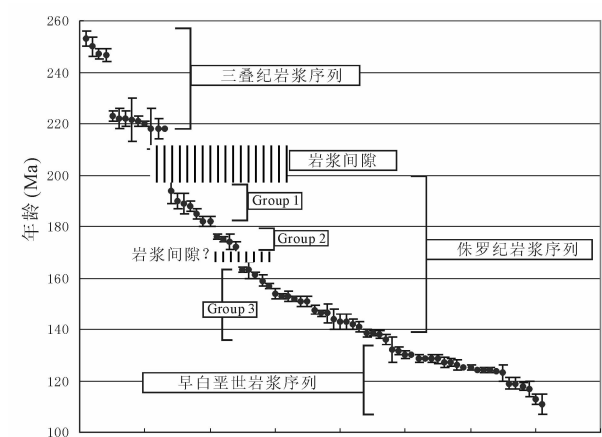


图 4 燕山—辽西带中生代侵入岩同位素年龄散点图

Fig. 4 Scatter plot of isotopic dating ages of Mesozoic intrusive rocks from the Yanshan-Liaoxi belt

数据来自: Davis et al. (2001); 韩宝福等 (2004); 牟保磊等 (2001); 毛德宝等 (2003); 罗镇宽等 (2001a, b, 2003, 2004); Luo et al. (2001); 刘翠等 (2004); 任康绪等 (2004); Li et al. (2003); 苗来成等 (2003); 吴福元等 (2006); Yan et al. (2000); 苏尚国等 (2006); 张宏等 (2002)

Data sources from: Davis et al. (2001); Luo et al. (2001a, b, c, 2003, 2004); Mu et al. (2001); Han et al. (2004); Mao et al. (2003); Liu C et al. (2004); Li Y et al. (2003); Miao et al. (2003); Ren K et al. (2004); Wu et al. (2006); Yan et al. (2000); Su et al. (2006); Zhang et al. (2002)

发展具有幕式和挤压-伸展交替演化特征。通过综合分析和对比华北地区侏罗纪地层、岩浆活动序列和构造变形事件,建立了侏罗纪大地构造发展的三阶段演化历史,从早到晚,阐述如下。

阶段 I: 早侏罗世早期挤压阶段 (205~191 Ma)。这一阶段显著特征是缺乏侵入岩浆活动,华北地块内部发生区域隆升和剥蚀,其代表性的标志是鄂尔多斯盆地和合肥盆地下侏罗统底部的侵蚀不整合面。侵入岩体仅零星出露于辽西地区。对这个阶段的构造变形特征认知较少,尽管在燕山地区识别出了发生在 180Ma 之前的构造挤压变形事件(郑亚东等, 2000; Davis et al. , 2001)。这个阶段板块运动背景包括:华北地块和华南地块沿着秦岭—大别—苏鲁构造带的持续陆内汇聚,北部西伯利亚板块与蒙古—华北板块的持续碰撞和汇聚(Yin and Nie, 1996; Chang, 1996)。

阶段 II: 早中侏罗世弱伸展 (190Ma~165±5 Ma)。该阶段构造相对比较平静,华北地块岩石圈处于弱伸展状态,表现为:① 沿燕—辽构造带和郯庐断裂带两侧,发生幔源岩浆侵入和基性到酸性火山喷

发;② 沿阴山—燕山构造带,发生强烈的地壳引张和伸展盆地的形成和发育;③ 华北地块整体发生沉降,形成一个统一的含煤拗陷,覆盖了现今的鄂尔多斯盆地、山西高地和华北平原。垂向差异抬升和下降控制着早中侏罗世地层的沉积序列。鄂尔多斯盆地及其周缘地区的正断层发育,它们继承了华北地块基底 WE 向构造,引张方向为 NS 向至 NNE—SSW 向(张岳桥等, 2007)。从区域角度来看,这个阶段沿阴山—燕山构造带的伸展作用较强,而沿秦岭—大别构造带的伸展作用较弱。早中侏罗世华北地块伸展作用的动力可能与三叠纪和早侏罗世地块碰撞作用后的板内应力调整有关,这个阶段的大地构造属性可看作是印支运动的后效(董树文等,2007)。

阶段 III:中晚侏罗世多向挤压阶段($165\text{Ma}\pm 5\sim 136\text{Ma}$)。自中侏罗世晚期,区域板块运动学发生重大调整,东亚构造体制发生重大转换,华北地块遭受多向挤压变形,早中侏罗世形成的大型统一含煤盆地发生肢解,华北地块构造发生东西分异,鄂尔多斯盆地由此形成。这个阶段,至少发生两幕主要的挤压变形事件(见表 1):① 早幕以上侏罗统砾岩层底部的角度不整合为标志(鄂尔多斯盆地西缘的芬芳河组,北京西山的髻髻山组或九龙山组,辽西地区的蓝旗组),这个地层的顶部年龄在 $156\sim 166\text{Ma}$ (季强等, 2004;刘建等,2006;孙立新等, 2007),而构造变形的起始时代应该更早,但晚于 170Ma 。在燕-辽构造带识别出的 $170\sim 165\text{Ma}$ 的岩浆作用间隙期,可能是对这一挤压变形事件的响应。事实上,沿着燕山构造带和胶辽地块 $160\sim 150\text{Ma}$ 的同构造“S”型花岗岩原地侵位,以及北京西山和燕—辽地区同期基性到酸性火山作用,均指示中侏罗世晚期华北地块东部一次重要的地壳增厚事件。② 晚幕挤压变形以鄂尔多斯盆地和阴山—燕山构造带下白垩统和上侏罗统之间的角度不整合为标志,发生时代在 $135\pm 1\text{Ma}$ (张家口火山岩底部)之前、 150Ma 之后。胶辽和徐淮地区 $143\sim 135\text{Ma}$ 之间短暂的岩浆活动间隙,可能是对这一挤压变形事件的响应。

这一阶段华北地区构造变形表现为不同方向的地壳挤压缩短(见图 5)。北缘阴山—燕山构造带受 NNW—SSE 向挤压而冲断活化,南缘秦岭—大别构造带受 NNE—SSW 向挤压而逆冲活化,鄂尔多斯盆地西缘受近 WE 向挤压而活化,山西台地受 NW—SE 向挤压发生断褶变形,郯庐断裂带发生左旋走滑活动,郯庐断裂带东侧胶辽块体的构造变形既有逆冲、又有左旋走滑,是 NW—SE 向至

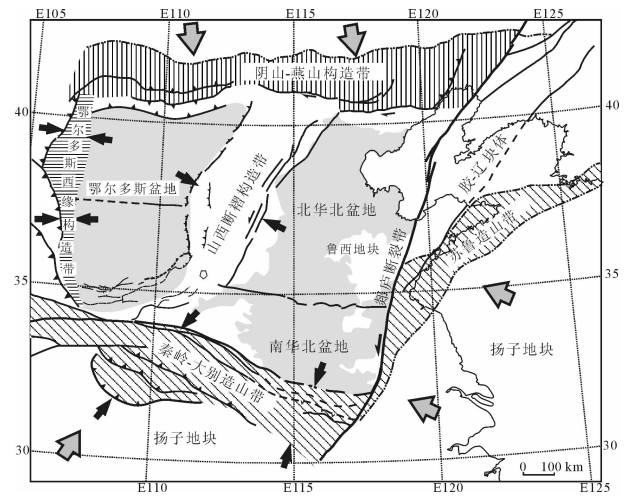


图 5 华北中晚侏罗世多向挤压构造简图

Fig. 5 Schematic map showing Middle to Late Jurassic multi-directed contractions in North China

NNCP—北华北盆地; SNCP—南华北盆地; LXB—鲁西地块;
TLFZ—郯庐断裂带

NNCP—northern North China Plain; SNCP—southern North China Plain; LXB—Luxi block; TLFZ—Tancheng-Lujiang fault zone

NW—SEE 向挤压结果。

5 讨论

5.1 中晚侏罗世地壳增厚作用、构造样式和应变分解

多方面的证据显示华北地块岩石圈地壳在中侏罗世经历了一次重要的增厚过程。① 在郯庐断裂带东侧的胶辽地块和郯庐断裂西侧的蚌埠隆起带、以及沿燕-辽构造带出现大量的地壳重熔型花岗岩,指示了中侏罗世末期华北地壳发生了重要的增厚作用;② 地球化学研究结果暗示,燕山构造带晚侏罗世髻髻山组或蓝旗组中的火山活动发生在挤压构造背景下(李伍平等, 2001, 2004)。路凤香等(2006)估算在髻髻山组火山喷发之前,地壳厚度达到 $40\sim 45\text{km}$;③ 此外,沿鄂尔多斯盆地西缘发育的晚侏罗世芬芳河组、沿燕辽构造带发育的土城子组或后城组、沿大青山发育的大青山组,均是对中晚侏罗世地壳挤压增厚事件的沉积响应。

华北地块中晚侏罗世挤压构造展示了具有显著对照的、东西分异的构造样式。一方面,山西台地的逆冲断层及其相关褶皱被解释为与上地壳内的滑脱构造有关,而下地壳没有卷入到缩短变形,可以将此看作是薄皮构造。这种薄皮构造样式同样适用于大青山逆冲推覆构造带。另一方面,位于郯庐断裂带

东侧的胶辽地区,大量晚侏罗世花岗岩的侵位指示了卷入下地壳的厚皮构造特征。在郯庐断裂带西侧蚌埠隆起带上,小规模片麻状 S 型花岗岩的侵位同样指示了该区卷入下地壳的厚皮构造变形特征。在山西台褶带中未见中晚侏罗世岩体侵入,这个事实说明,太行山东缘——一条著名的重力梯度带,成为了东部地区卷入下地壳的厚皮构造向西部地区卷入上地壳的薄皮构造的转换带。这种构造变形样式进一步指示,中晚侏罗世的地壳增厚作用主要发生在位于太行山以东的华北东部地区。

两个纬向构造带(阴山—燕山构造带和秦岭—大别构造带)侏罗纪变形样式同样存在东西分异的特点。例如,卷入下地壳的燕山—辽西构造带,晚侏罗世岩浆作用十分活跃,而仅卷入上地壳的阴山构造带,没有晚侏罗世岩浆活动。

相应地,中晚侏罗世的地壳变形在时间和空间上分解为逆冲作用和走滑活动。通常的解释是,近 NS 向地壳挤压和缩短变形主要发生在两个纬向构造带,而 NW-SE 向挤压和地壳缩短变形主要发生在华北地块的广大地区,包括了山西高地、华北平原北部和胶辽块体。特别地,郯庐断裂带东侧的胶辽块体以逆冲和走滑共同作用为主要特征,山西高地以逆冲和断层相关褶皱变形为主,北华北平原区则以左旋走滑为主。这一应变分解型式指示了华北地块在中晚侏罗世时期多向挤压的大地构造背景。

5.2 华北地区中晚侏罗世多向挤压变形的板块动力学背景

华北地区中晚侏罗世多向挤压变形的识别具有重要的大地构造意义。自从翁文灏(Wong, 1927, 1929)提出燕山运动之后,华北地块的中晚侏罗世构造事件研究已经有 80 余年的历史。赵越等(1994, 2004)强调,燕-辽和北京西山地区的中晚侏罗世构造热事件,记录了古亚洲和特提斯构造域向古太平洋构造域的转换过程,起始于 170 Ma,终止于 135 ± 1 Ma。最近十年,有关中国东部燕山运动的本质和晚中生代构造体制转换过程问题成为中国地学界讨论的热点主题(Zhai et al., 2004; 翟明国等, 2004; 吴福元等 2000, 2003; 周新华, 2006)。在这些讨论中,人们往往聚焦于华北岩石圈减薄和克拉通破坏问题,却忽视了岩石圈减薄之前发生的地壳增厚作用。笔者认为,正是中国东部中晚侏罗世时期的多向挤压和地壳增厚,才诱发了白垩纪时期的岩石圈伸展和巨量减薄。

事实上,侏罗纪是东亚大陆构造体制转折和板

内变形的重要阶段,周邻板块向东亚大陆同时汇聚导致了华北地区中晚侏罗世多向挤压变形和不同构造体系复活或新生。这个时期板块活动边界有 3 个:亚洲东部的古太平洋俯冲消减带,位于羌塘和拉萨地块之间的碰撞俯冲带,位于西伯利亚地块和蒙古—华北板块之间的碰撞带。华北地块远离这些活动板块边界,但处于这些活动板块边界的中心部位。其中因古太平洋板块向亚洲大陆之下俯冲作用形成的东亚活动增生陆缘的出现深刻地影响着华北地区中晚侏罗世变形格局。三大板块——北部西伯利亚板块、东部古太平洋板块和西部拉萨地块,同时向东亚大陆碰撞汇聚产生的远程效应主导了华北地区中晚侏罗世多向挤压变形和应变分解。

古蒙古—鄂霍茨克洋在中晚侏罗世开始闭合(Parfenov and Natal'in, 1986; Zhao X et al., 1990; Zorin et al., 1995; Nie and Rowley, 1994; Van der Voo et al., 1999),西伯利亚板块和蒙古—华北板块之间的碰撞汇聚持续到晚侏罗世,其产生的远程效应激活了古老的阴山—燕山构造带,使其发生强烈的挤压逆冲变形。北山—戈壁地区发现的巨大逆冲推覆岩席(Zheng et al, 1996),是远离板块边界的蒙古—华北板块向南运动的最好例子。在青藏地区,拉萨地块和羌塘地体沿着班公—怒江缝合带在晚侏罗世的碰撞(约 160 Ma),导致羌塘盆地下白垩统和上侏罗统地层之间的不整合(贾承造等, 2005)。刘永江等(Liu et al, 2003)报道了沿阿尔金断裂东段的变形花岗片麻岩的⁴⁰Ar-³⁹Ar 年龄(178.4~137.5 Ma),记录了这期构造-热事件的年代。Wang Y 等(2005)通过⁴⁰Ar-³⁹Ar 测试,获得了 165~160Ma 阿尔金断裂左旋走滑活动的冷却年龄,与班公湖—怒江缝合带的碰撞时代一致。这个块体碰撞导致了海原断裂带北部、向东变宽、呈楔形展布的阿拉善块体和陇西块体的向东挤出,进而导致了鄂尔多斯盆地西缘基底薄弱带的逆冲活化(张岳桥等, 2007)。在亚洲东缘沿着太平洋边缘,古太平洋板块在中侏罗世(170Ma 左右)向亚洲大陆之下发生俯冲作用(Isozaki and Nishimura, 1989; Isozaki, 1997; Maruyama et al., 1997)。Isozaki 等(1990)相信,大部分日本俯冲增生楔的形成时代可以追溯到 160 Ma 之前的板块俯冲。太平洋板块在中晚侏罗世初期可能以低角度俯冲于亚洲大陆之下,引起了大陆地壳的增厚以及华北东部地区逆冲和左旋走滑的应变分解。华北地壳增厚机制值得今后进一步研究。

6 结论

侏罗纪是东亚大陆板内构造发展的重要时期。华北地区侏罗纪构造可以划分为三个演化阶段。早中侏罗世(190~170 Ma)弱伸展作用阶段代表了一个构造转换阶段,从早期以华北—华南地块沿秦岭—大别—苏鲁构造带碰撞和西伯利亚与蒙古地块的陆-陆碰撞为特征的古亚洲构造域,向晚期(从165±5 Ma 至 136 Ma)以多向挤压和强烈的陆内变形为特征的燕山运动阶段转变。燕山运动构造呈幕式发展,挤压变形在空间上发生分解,其动力背景与周邻 3 大板块(北部的西伯利亚板块、东部的古太平洋板块和西部的特提斯板块)同时向东亚大陆汇聚产生的远程效应有关。华北地块正处于这样一个多向板块汇聚背景的核心部位,其东部地区的地壳和岩石圈发生显著增厚,并成为诱发中国东部早白垩世岩石圈巨量减薄的关键因素。郯庐左旋走滑断裂系正是在这个燕山运动多向挤压阶段中发展的。

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Jurassic Tectonics of North China: a Synthetic View

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Abstract

This paper gives a synthetic view on Jurassic tectonics of North China, with an attempt to propose a framework for stepwise tectonic evolution history. Jurassic sedimentation, deformation and magmatism in North China have been divided into three stages. The earliest Jurassic is marked by a magmatism gap in 205~190 Ma and regional uplift, which are considered to be the continuation of the "Indosinian Movement" tectonic period characterized by continent-continent collision between the NCB and the SCB. Early to Middle Jurassic (190~170 Ma) was dominated by weak lithospheric extension expressed by mantle derived plutonism and volcanism along the Yanshan belt and alongside the Tan—Lu fault zone, normal faulting and graben formation along the Yinshan-Yanshan tectonic belt, depression and resumption of coal-bearing sedimentation in vast region of the NCB. Middle to Late Jurassic stage (from 165 ± 5 Ma to 136 Ma) was dominated by intensive intraplate deformation resulting from multi-directed compressions. At least two major deformation events have been identified; one is marked by stratigraphic unconformity beneath thick Upper Jurassic molassic series in the foreland zones of the western Ordos thrust and fold belt and along the Yinshan-Yanshan belt and predated to 160 Ma; the other one is indicated by stratigraphic unconformity at the base of Lower Cretaceous and predated to 135 Ma. During this last stage, two latitudinal tectonic belts, the Yinshan-Yanshan one to north and the Qinling—Dabie one to south, and the western margin of the Ordos basin were all activated by thrusting; the NCB itself was deformed by NE to NNE-trending structural system involving thrusting, associated folding and sinistral strike-slip faulting, which were spatially partitioned. Foliated S-type granitic plutons aged 160~150 Ma were massively emplaced in the Jiao-Liao region east of the Tan-Lu fault zone and indicate important crustal thickening in the eastern part of North China. Jurassic deformation patterns, different tectonic systems and multi-directed contractions in North China recorded far-field effects of synchronous convergences, toward the east Asian continent, of three different plates, the Siberian in north, the paleo Pacific in east and the Lhasa Block in southwest. This Middle to Late Jurassic intraplate orogenesis and pervasive shortening deformation preceded lithospheric attenuation and thinning in East China, which most possibly started by the Earliest Cretaceous around 135 Ma.

Key words: Jurassic tectonics; intraplate deformation; North China; multi-directional compressions; convergent tectonic system