

“秦岭-大别-苏鲁”造山带中“古特提斯缝合带”的连接

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内容提要:中国大陆西北部的“古特提斯缝合带”如何与东面的“秦岭-大别-苏鲁”造山带连接,是涉及中国大陆中部构造格架的关键问题之一。南秦岭造山带中的古特提斯蛇绿岩带和东秦岭-桐柏-大别-苏鲁造山带中三叠纪高压-超高压变质带的对比,以及一条位于两者之间的220~204 Ma的大型左行走滑剪切带的存在,提供了它们之间关系的新的视角,为此,我们提出南秦岭的勉略蛇绿岩带向东通过宁陕-湘河大型左行走滑剪切带,和大陆俯冲与深俯冲造成的“耀岭河-桐柏-大别-苏鲁”高压-超高压变质带北缘连接,构成“秦岭-大别-苏鲁”造山带中的古特提斯缝合带新模式。沿着这条边界,南秦岭构造单元可以分为南部的南秦岭被动陆缘单元和北部的南秦岭主动陆缘单元,后者向东的延伸由于南、北板块之间三叠纪的剪切碰撞而尖灭。

关键词:秦岭-大别-苏鲁造山带;古特提斯蛇绿岩带;高压-超高压变质带;宁陕-湘河左行走滑剪切带

1 问题的提出

东西向延伸、长约2000 km的秦岭-大别-苏鲁造山带是中国大陆的脊梁。晚元古代以来,从冈瓦纳大陆分离的南、北中国板块,经过始特提斯洋盆和古特提斯洋盆的演化以及板块多次离散、汇聚和碰撞,形成显生宙以来始特提斯和古特提斯为主体的复合构造格架,以及早古生代和印支期为主体的秦岭-大别-苏鲁复合造山系(Mattauer et al., 1985; Hsu et al., 1987; 许志琴等, 1988; 张国伟, 1988)。研究表明,大别-苏鲁造山带以大面积出露的三叠纪含柯石英和金刚石的高压-超高压变质带为特征(Xu, 1987; Wang et al., 1989; Okay et al., 1989; 杨建军, 1991; Xu et al., 1992; 王清晨等, 1992; 张希道等, 1992, 1999; Wang et al., 1993; Zhang et al., 1995, 2000; Zhang and Liou, 1997; Kato et al., 1997; Xu and Su, 1997; 叶凯等, 1999; 郑常青等, 1999), 250~200 Ma期间,南中国板块的表壳岩曾俯冲和深俯冲在北中国板块之下达到地幔的深度(>100 km),而后又折返至地表(Li et al., 1993,

2000; Liu et al., 2001)。

秦岭造山带的研究已经有半个多世纪(姜春发等, 1963, 1979; 黄汲清等, 1977; 李春昱等, 1978, 1982; 任纪舜等, 1980, 1990, 1991; 王鸿桢等, 1982; Mattauer et al., 1985; Hsu et al., 1987; 许志琴等, 1988; 张国伟等, 1988, 1991, 1995, 1996, 1989, 2001; 张本仁, 1990; Ratschbacher, L. et al., 2003., 王宗起等, 2009; Dong, Y. P., 2011; Wu YB and Zheng YF, 2013; Bader T. et al., 2013a, b)。由于秦岭-大别-苏鲁造山带构造的复杂性、多期性、复合性、叠置性和穿时性,对于造山带的基本构造格架一直存在分歧和争议,其中,“古特提斯缝合带”的位置和链接,被认为是关键科学问题之一。

秦岭-大别-苏鲁造山带中是否存在古特提斯缝合带?如何连接?至今争议不断。李春昱(1978, 1982)曾认为位于西秦岭的勉(县)略(阳)缝合带是一条印支板块缝合带,向西链接东昆仑南缘的阿尼玛卿缝合带;张国伟等(1995, 1996, 2001)提出勉略带是一条由岛弧火山岩、洋脊蛇绿岩残片和混杂堆积组成蛇绿构造混杂带,形成时代220~240 Ma,向

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东经南秦岭和四川盆地之间的大巴山前陆弧形构造带延至大别山南缘。

笔者通过秦岭-桐柏地区的深入野外调查,结合前人研究成果,提出南秦岭古特提斯缝合带的另一连接方式:即位于西秦岭的勉略古特提斯缝合带向东经 NEE—SWW 向的“洋县-宁陕-湘河”走滑剪切带的左行错位,链接南秦岭北部的耀岭河高压变质带,向东延至桐柏高压变质带、红安-大别高压—超高压变质带,经大型郯庐走滑断裂的左行错位,抵达 NE—SW 向的苏鲁高压—超高压变质带(图 1)。

2 “秦岭-大别-苏鲁”造山带中“古特提斯缝合带”连接的依据

2.1 勉略蛇绿混杂岩带与白水江增生楔

西秦岭勉略蛇绿混杂岩带的厘定(张国伟等,

2001, 2003, 2004)开启了东昆仑南缘阿尼玛卿古特提斯缝合带(308 Ma, 杨经绥等, 2004)深入秦岭造山带的大门。张国伟等(1995, 1996, 2001, 2003, 2004)认为勉略带代表泥盆纪中期开始至三叠纪的以深水浊积岩为特征的古特提斯有限洋盆, 并伴随 241~220 Ma 俯冲岛弧的形成(李曙光等, 1996)和晚三叠世古特提斯洋盆闭合以及印支碰撞造山。勉略蛇绿岩带北部大面积分布了白水江群大堡组(原定奥陶纪)泥岩和粉砂岩中夹有灰岩、硅质岩、玄武岩和凝灰岩块体, 最近由于在灰岩中发现中泥盆世的牙形石和珊瑚化石(王涛等, 2011), 表明白水江群有可能为晚古生代形成的增生楔。这样, 勉略蛇绿混杂岩带与白水江增生楔一起构成西秦岭的古特提斯板块边界。

另外, 在勉略带中还发现具有岛弧性质的基性

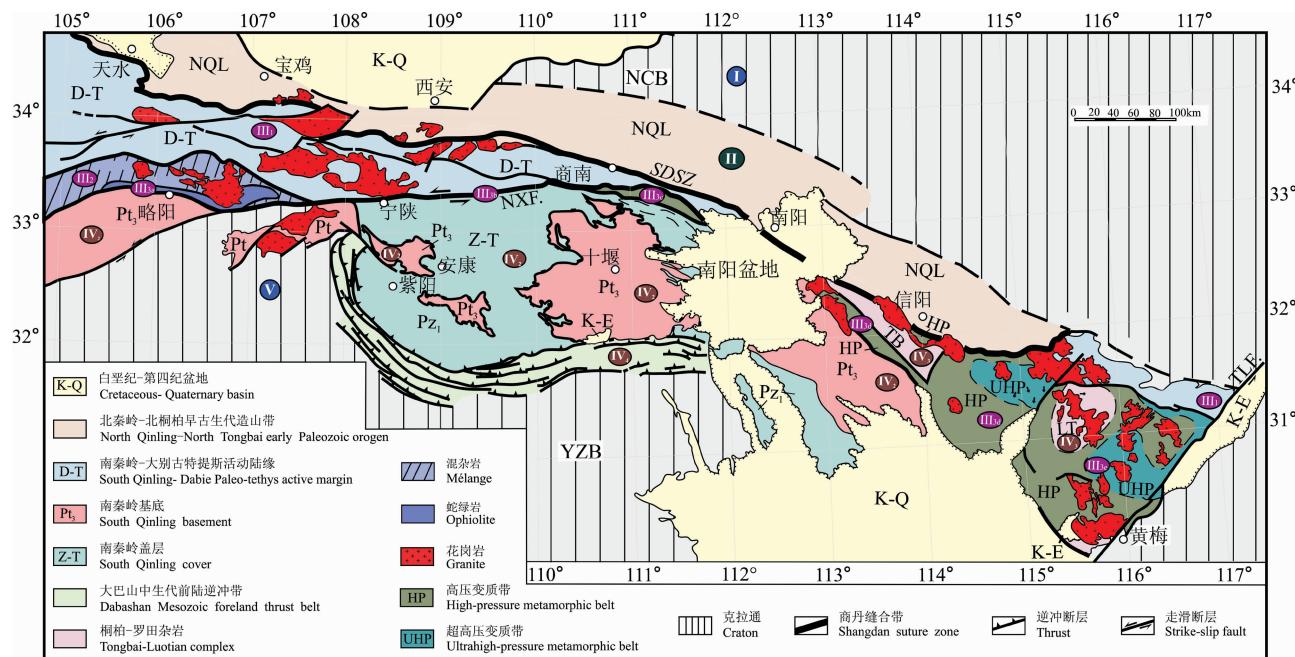


图 1 秦岭-桐柏-大别-苏鲁造山带中的古特提斯构造格架图

Fig. 1 Paleo-Tethys tectonic framework of the Qinling-Tongbai-Dabie-Sulu Orogen

I—北中国陆块(NCB); II—北秦岭-北桐柏早古生代造山带(NQL); III + IV—南秦岭-大别三叠纪造山带: III₁—南秦岭-大别古特提斯活动陆缘(D-T); III₂—白水江古特提斯增生混杂带; III_{3a}—勉略古特提斯蛇绿混杂岩带; III_{3b}—宁陕-湘河左行走滑剪切带; III_{3c}—耀岭河高压变质带; III_{3d}—桐柏-大别高压变质带(HP); III_{3e}—大别超高压变质带(UHP); IV—南秦岭-南桐柏被动陆缘: IV₁—南秦岭盖层(Z-T、Pz₁); IV₂—南秦岭基底(Pt₃); IV₃—桐柏-罗田中生代杂岩(TB、LT); IV₄—大巴山中生代前陆逆冲带; V—扬子陆块(YBZ); SDSZ—商丹缝合带; NXF—宁陕-湘河左行走滑断裂; TLF—郯庐断裂

I—North China Block; II—North Qinling Early Paleozoic Orogen; III + IV—South Qinling Triassic Orogen: III₁—South Qinling-Dabie Paleo-Tethys active continental margin; III₂—Baishuijiang Paleo-Tethys accretional belt; III_{3a}—Mianlue Paleo-Tethys ophiolitic mélange zone; III_{3b}—Ningshan-Xianghe left-lateral strike-slip shear zone; III_{3c}—Yaolinhe HP metamorphic belt; III_{3d}—Tongbai-Dabie HP metamorphic belt; III_{3e}—Dabie UHP metamorphic belt; IV—South Qinling-South Tongbai passive continental margin; IV₁—South Qinling covers (Z-T); IV₂—South Qinling basement; IV₃—Tongbai-Luotian Mesozoic complex; IV₄—Dabaishan foreland thrust belt; V—Yangtze Block. SDSZ—Shangdan Suture zone; TLF—Tanlu fault

麻粒岩(李三忠等,2000),其原岩形成时代为477 Ma,是在早于215 Ma之前俯冲至地壳60 km深度,遭受高压榴辉岩相和高压麻粒岩相变质作用,214 Ma开始折返成角闪岩相(梁莎等,2013)。

2.2 耀岭河-桐柏-大别-苏鲁(超)高压变质带

位于商丹缝合带以南的泥盆—石炭纪刘岭群复理石单元(或增生楔)(Mattauer et al., 1985; 许志琴等, 1988; You et al., 1993; Ratschbacher et al., 2003, 2006)之东南出露了前寒武纪变质基底—耀岭河群变基性火山岩系。由于赵川地区蓝闪石类矿物的发现,使之成为高绿片岩-蓝片岩变质岩相的高压变质带、钠闪石的Ar-Ar年龄为 232 ± 5 Ma,多硅白云母年龄为 216 ± 7 Ma,曾被认为是南秦岭陆内深层滑脱构造的产物(Mattauer et al., 1985; Xu, 1987; 许志琴等, 1988)。南秦岭以东的桐柏高压变质带位于桐柏杂岩的南、北两侧,高压榴辉岩的峰期变质年龄为255~256 Ma,冷却年龄为238~217 Ma(Liu et al., 2008, 2010; Cheng et al., 2011; 刘晓春等, 2013)。向东的红安高压—超高压变质带的超高压峰期变质年龄被限定于239~226 Ma之间(Wu et al., 2008; Zhou et al., 2011),折返冷却年龄为216~212 Ma(Liu et al., 2004; Jahn et al., 2005; Wu et al., 2011; Li et al., 2011)。再向东的大别-苏鲁高压-超高压变质带中,高压岩片的峰期变质年龄为260~245 Ma,折返年龄为253~210 Ma(Eide et al., 1994; Webb et al., 1999; Liu Y C et al., 2000; Xu et al., 2006, 2009; Liu et al., 2006);超高压岩片的峰期变质年龄为240~225 Ma,折返年龄为220~200 Ma(Eide et al., 1994; Webb et al., 1999; Hacker et al., 2000; Ayers et al., 2002; Li et al., 2000, 2003; Liu and Xu, 2004; 许志琴等, 2005; 李曙光等, 2005; Xu et al., 2006; Wu et al., 2006; Liu et al., 2005, 2006, 2007, 2009; Zhang et al., 2007; Liu Y C et al., 2007; Liu and Liou, 2011)。

上述结果表明,耀岭河、桐柏、大别和苏鲁的高压变质带,以及大别-苏鲁的超高压变质带均具有“形成-折返”时间上的一致性和空间上的关联性。因此,耀岭河地区高压变质带可与桐柏-大别-苏鲁地区的(超)高压变质带连接(Wu and Zhen, 2013)。尽管耀岭河-桐柏高压变质带呈带状展布、大别-苏鲁(超)高压变质带呈面状分布,但均反映东秦岭以东地区的南中国板块,沿着(超)高压变质带北缘的板块边界,俯冲或深俯冲于北中国板块之下、继后折

返至地表的重大构造事件。因此,耀岭河-桐柏-大别-苏鲁(超)高压变质带的北缘实际上应为南、北中国板块之间的缝合带位置。

2.3 宁陕-湘河左行走滑断裂

勉略古特提斯蛇绿混杂岩带和三叠纪耀岭河高压变质带之间可以通过长达200 km的近东西向的宁陕-湘河左行走滑剪切带链接。以勉略-宁陕-耀岭河为界,可以将南秦岭的古特提斯构造单元划分成两部分:南秦岭北部的活动陆缘地体和南秦岭南部的被动陆缘地体。

南秦岭北部活动陆缘地体主要由西成、凤太和山柞盆地内的泥盆—二叠系和阿尼玛卿缝合带北侧西秦岭三叠系以及相关的岛弧钙碱性花岗岩共同构成。其中泥盆—二叠系为一套滨浅海—深海相沉积组合,主要为一套成熟度较低的杂砂岩夹砂屑灰岩或者生物碎屑灰岩,在垂直于造山带走向方向该套组合具有向南水体逐渐加深且地层时代逐渐变年轻特征;物源分析、砂岩碎屑组成和地球化学成分综合分析表明其形成于活动大陆边缘弧前环境(闫臻等, 2002, 2007; 陈义兵等, 2010; 吴树宽等, 2012; Yan et al., 2006a, 2006b, 2012; 王元元等, 2014)。同时该套组合被不同断裂分割成一系列次级沉积盆地,呈现出弧前背驮式盆地特征(王宗起等, 2002, 2009)。三叠系主要分布于南秦岭西段的西秦岭楔,由下—中三叠统隆务河群、古浪堤组、留凤关群和上三叠统鄂拉山群构成,其中下三叠统主要是一套大陆斜坡-深海盆地相沉积组合并夹有丰富的安山岩和英安岩,而上三叠统主要由安山岩、英安岩、流纹岩和中酸性火山碎屑岩组成,它们与同时期的岛弧钙碱性闪长岩、花岗闪长岩、英云闪长岩、花岗岩和辉长岩(张宏飞等, 2006; Guo et al., 2012; Li et al., 2015)共同构筑了三叠纪活动大陆边缘(闫臻等, 2008, 2012; Yan et al., 2014)。这些泥盆—三叠系活动大陆边缘沉积组合向东与桐柏-大别造山带中的泥盆系南湾群和龟山群相对应(Liu et al., 2013)。

比较而言,南秦岭被动陆缘地体主要由新元古代火山-沉积岩组合的基底与寒武—三叠纪大陆边缘斜坡相-深海相沉积盖层共同构成。红椿坝断裂和安康断裂将该被动陆缘地体分割成若干个次级盆地沉积组合序列。这些沉积组合自南而北表现为寒武—三叠系连续沉积,由浅水灰岩、白云岩和砂岩为主的潮坪和内陆棚环境向北部逐渐过渡为以板岩、泥灰岩、夹砂岩、薄层灰岩外陆棚环境和以浊积岩为

特征的复理石构成。奥陶—志留纪时期,在高滩-镇坪和紫阳-竹溪盆地内发育有丰富的碱性火山岩及相关火山沉积组合,显示了被动陆缘火山裂谷作用的特点;丰富的底栖生物化石和笔石化石以及向南砂岩增多变粗、厚度变大并富含火山碎屑岩物质特征,也表明南秦岭被动陆缘在奥陶—志留纪时期经历了强烈的伸展作用。

观察和研究表明,“宁陕-湘河”左行走滑断裂(图 1)是由 2~2.5 km 宽的韧性左行剪切带组成,在湘河一带为糜棱岩化的晚元古代的花岗片麻岩,在洋县宁陕一带由志留—泥盆纪千糜岩化的石英片岩和强烈变形带组成,面理 E—W 向近直立,拉伸线理近水平,并含同构造的花岗岩脉(图 2b)。剪切应变显示明显的高温左行走滑特征。该剪切带中段可见古生代—三叠世地层的拖曳构造,显示左行走滑特征(图 2a),推测该断裂形成于晚三叠世—早侏罗世,为一条南秦岭古特提斯缝合带的转换带。

2.4 古特提斯缝合带的组成和东、西差异

代表西部勉略蛇绿岩带的侵位和印支碰撞造山发生在晚三叠—早侏罗世(张国伟等,1995,2004),而东部以“桐柏-大别-苏鲁高压-超高压变质带”为标志的南、北中国板块碰撞和深俯冲的时限应在早三叠世之前(>250 Ma),两者相差约 50 Ma。南、北中国板块之间汇聚和碰撞时限的东早西晚,产生剪切(点)碰撞的模式(董树文等,2002; Xu et al., 2009)。这一俯冲-碰撞的剪切模式使西部勉略缝合带呈线型展布,中部的桐柏高压变质带折返隆升后呈带状分布在桐柏变质-深熔杂岩体(140~110 Ma)的两侧(Liu et al., 2010),而东部“桐柏-大别-苏

鲁”高压-超高压变质带组成的俯冲板片经历大面积向南折返的长期效应。地震反射剖面揭示苏鲁高压-超高压变质带呈穹窿式的面型叠置岩片,覆盖在南中国板块之上(Xu et al., 2009)。在大别超高压变质带之下裸露了早白垩世的罗田变质-深熔杂岩(Xu, 1987),与桐柏杂岩形成时代和成因相同。

与勉略古特提斯洋盆相伴随的俯冲岛弧(241~220 Ma)(李曙光等,1996)和与东昆仑-阿尼玛卿古特提斯洋盆相伴随的布尔汉布达-阿尼玛卿俯冲岛弧(P-T)一起,构成近乎连续的古特提斯蛇绿岩带和岛弧岩浆带;而在东部桐柏-大别-苏鲁高压-超高压变质带组成的缝合带北侧(北中国板块活动陆缘一侧),至今没有发现古特提斯蛇绿岩带和岩浆岛弧带。但是,在含有南中国板块基底的南苏鲁和含有北中国板块基底的北苏鲁地体中,同样发育三叠纪的高压-超高压变质带,因而笔者(许志琴等,2007)曾提出板块边界两侧地体双俯冲(深俯冲)而造成的“俯冲剥蚀模式”(图 3),用以解释东部汇聚边界蛇绿岩和岛弧消失的原因,以及为什么南秦岭活动陆缘地体西宽东窄,乃至向东湮没,推测其可能已经成为北中国板块南缘早古生代造山地体之下的俯冲板片。

2.5 缝合带中的古老基底年龄

近年来,对勉略带中相关的火山岩和辉长岩的 SHRIMP 和 LA-ICP-MS 锆石测年,获得新元古代的岩浆锆石年龄信息(闫全人等,2007),引起人们对勉略蛇绿岩形成时代的争议。笔者认为,代表古特提斯洋盆残片的蛇绿岩和蛇绿混杂岩带,在俯冲和折返的过程中,携带代表被动陆缘早期演化的物质



图 2 宁陕左行走滑剪切带的野外照片

Fig. 2 Field photographs of the Ningshan sinistral strike-slip shear zone

(a)—宁陕剪切带中同构造花岗质脉的左行走滑运动指向;(b)—宁陕剪切带中 E—W 向直立面理和水平拉伸线理

(a)—Syntectonic granitic veins show the sinistral shear sense; (b)—The sub-vertical foliation and horizontal lineation in the Ningshan shear zone



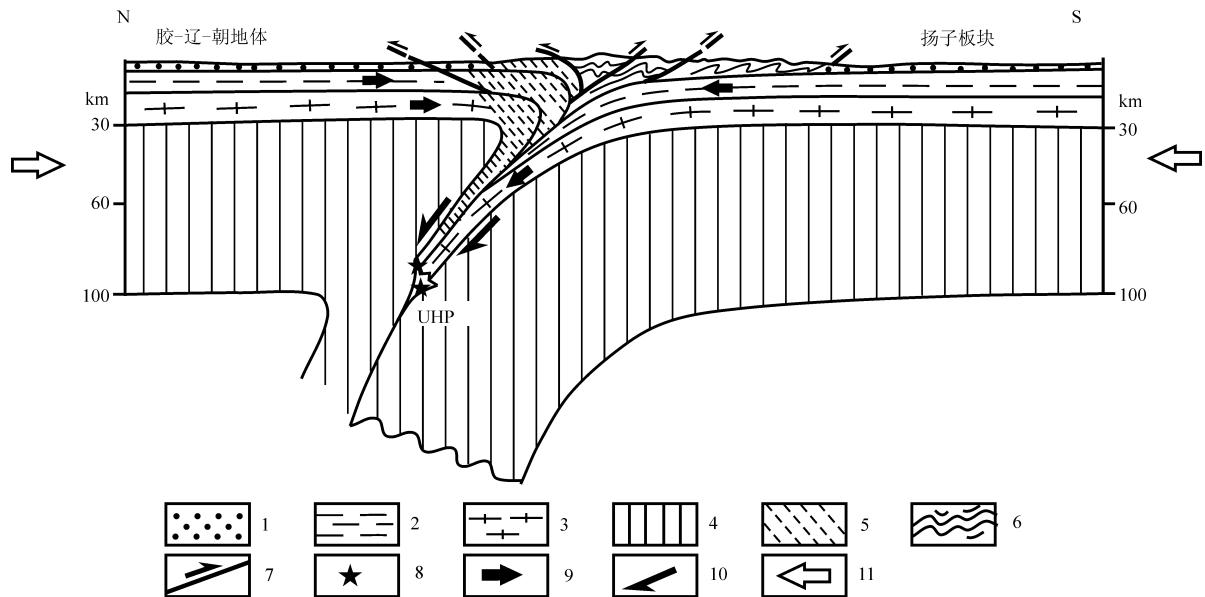


图3 苏鲁高压—超高压变质带形成的陆-陆碰撞俯冲剥蚀模式(据许志琴,2007)

Fig. 3 Deep subduction erosion model of continental-continental collision

for the formation of the Sulu HP-UHP metamorphic terrane (after Xu et al., 2007)

1—沉积层;2—上地壳;3—下地壳;4—岩石圈地幔;5—俯冲剥蚀板片;6—褶皱盖层;7—逆冲断裂;

8—超高压变质带;9—壳内滑动方向;10—板块俯冲方向;11—板块汇聚方向

1—Sediment; 2—upper crust; 3—lower crust; 4—lithosphere mantle; 5—subduction erosion slab; 6—folding covers; 7—thrust;

8—UHP metamorphic belt; 9—intercrustal sliding direction; 10—plate subduction vector; 11—plate convergence vector

返回浅部是完全可能的,而在勉略带南侧被动陆缘中只有新元古代具有岩浆事件的记录。同样,在耀岭河-桐柏高压变质带和大别-苏鲁(超)高压变质带中,发现来自古老基底的岩浆锆石年龄:如耀岭河高压变质岩的原岩是由一套新元古代岛弧和裂谷型火山沉积序列组成,其中岛弧型基性熔岩在 847 ± 8 Ma结晶,火山岩形成在 731 ± 11 Ma 和 768 ± 7 Ma之间,裂谷型熔岩形成在 $680 \sim 650$ Ma之间(Zhu et al., 2014);桐柏高压榴辉岩的原岩年龄为 $933 \sim 742$ Ma (Liu et al., 2008, 2010; Kröner et al., 1993);大别(超)高压变质岩的原岩年龄为 $798 \sim 684$ Ma (Ames et al., 1993, 1996; Rowley et al., 1997; Xue et al., 1997; Hacker et al., 1998, 2000; Liu et al., 2003, 2004; Zheng et al., 2006)。而东部苏鲁(超)高压变质带同时出露在位于南中国陆块一侧的南苏鲁和北中国陆块一侧的北苏鲁(即胶辽地块),在南苏鲁地区(超)高压变质岩石中保存 $780 \sim 700$ Ma 花岗岩浆和基性岩浆记录,为罗迪尼亞超大陆形成的双峰式岩浆裂解事件的产物(许志琴等,2006),而在北苏鲁的(超)高压变质岩石中存在北中国古老基底变质事件的年龄(> 2400 Ma, $1780 \sim 1821$ Ma)(许志琴等,2006)。因此,在缝合带和

(超)高压变质带中保留早期事件记录是可以理解的。

3 初步认识

(1) 南秦岭西部的勉略古特提斯蛇绿岩带向西与东昆仑的阿尼玛卿蛇绿岩带相接,向东通过宁陕-湘河大型左行走滑剪切带,和大陆俯冲与深俯冲造成的“耀岭河-桐柏-大别-苏鲁”高压—超高压变质带北缘连接。

(2) 确定的勉略-耀岭河-桐柏-大别-苏鲁缝合带分割了古特提斯时期南秦岭北部的活动陆缘和南秦岭南部的被动陆缘。

上述意见愿与同仁们商讨,不当之处,恭请批评指正。

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A Connection between of the Paleo-Tethys Suture Zone in the Qinling-Dabie-Sulu Orogenic Belt

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Abstract

How the "Paleo-Tethyan suture zone" of NW China is connected with the "Qinling-Dabie-Sulu orogen" to the east is one of the key issue in the tectonic framework of central China. Similarities between the Paleo-Tethys ophiolite zone in South Qinling and the Triassic HP-UHP metamorphic belt in the East Qinling - Tongbai-Dabie-Sulu orogen, as well the presence of a 214~200 Ma, large-scale, left-lateral shear zone between them, provide new insights into their relationship. Thus, we propose a new model in which the Mianlue ophiolitic belt (Zhang G W et al., 2001) is connected with the northern edge of the "Yaolinghe-Tongbai-Dabie-Sulu belt to the east by the Ningshan-Xianghe left-lateral shear zone. This shear zone divides the South Qinling terrane into the South Qinling passive continental margin unit to the south and the South Qinling active unit to the north, which pinches out to the east due to shearing between the NCB and SCB during the Triassic

Key words: Qinling-Dabie-Sulu Orogen; Paleo-Tethys Suture Zone; High-pressure and ultrahigh-pressure metamorphic belt; Ningshan-Xianghe strike-slip shear zone